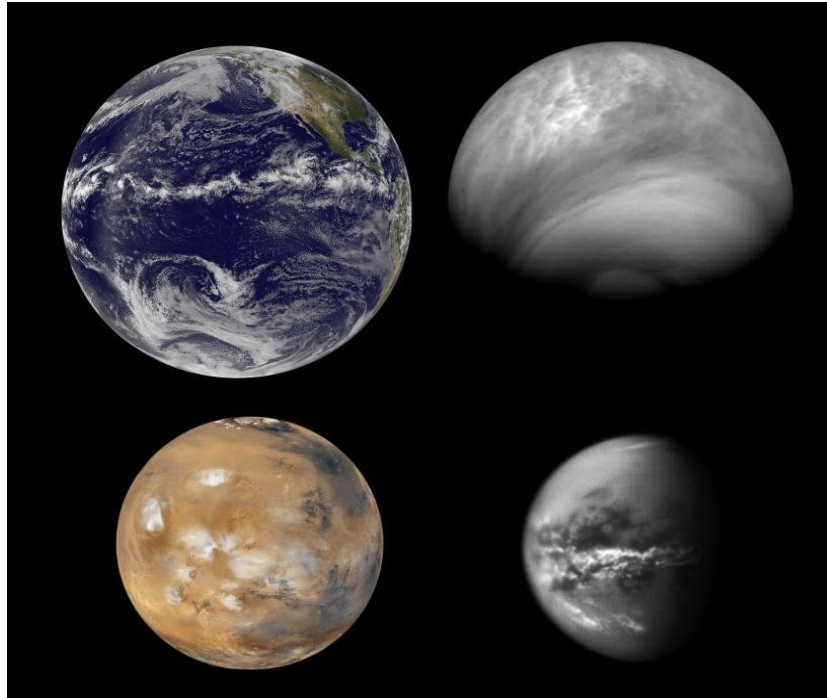


# PLANETARY ATMOSPHERES

## 4. Global Climate Modeling



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# PLANETARY ATMOSPHERES

## Global Climate Modeling

- Virtual planets
- Different models for different scales
- Successes, and lessons from failure

Forget F. and Lebonnois S., « Global climate models of the terrestrial planets »  
*Comparative Climatology of the Terrestrial Planets*, S.J. Maxwell et al. Eds.  
University of Arizona, Tuscon, 2013

# PLANETARY ATMOSPHERES

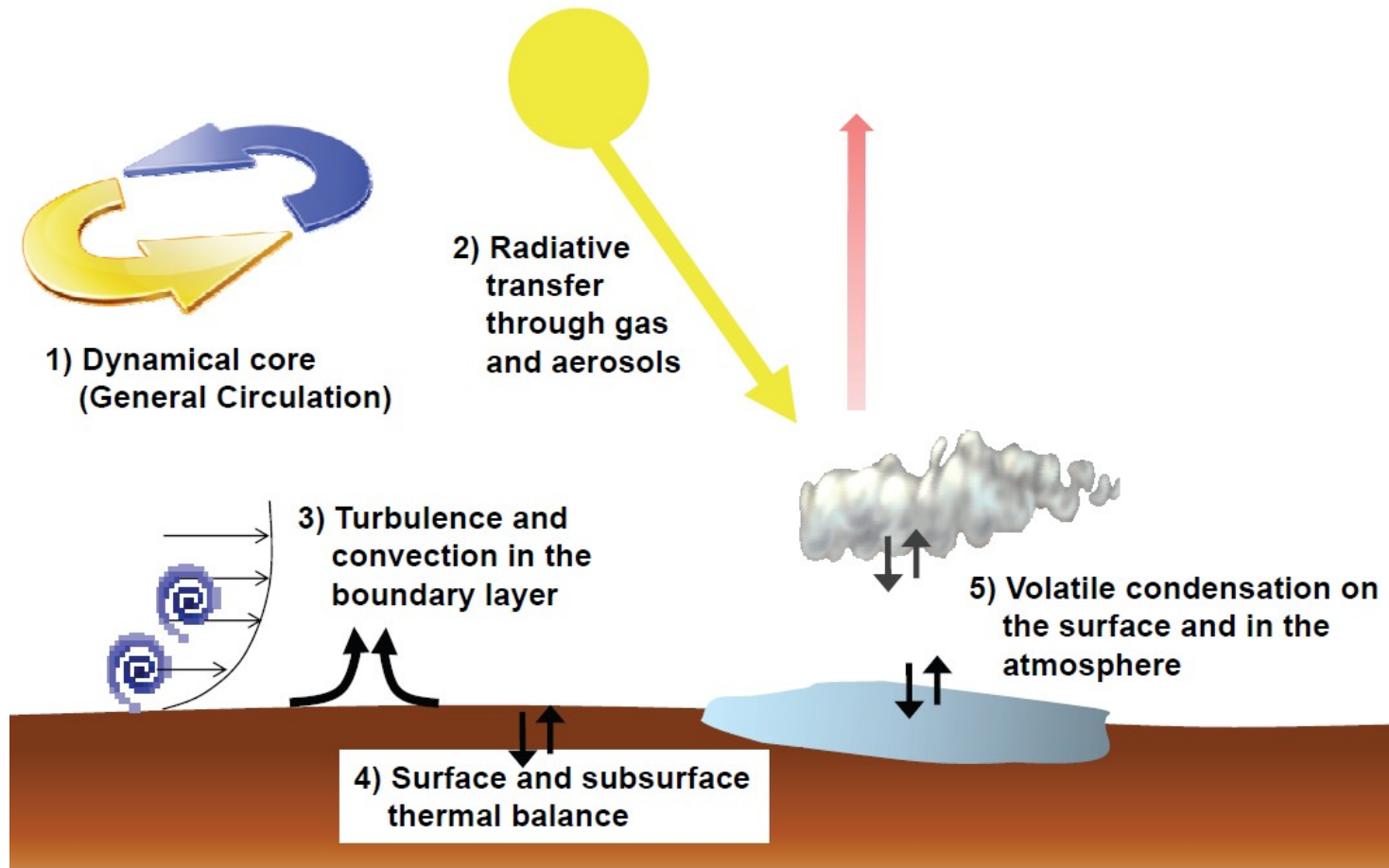
## Global Climate Modeling

- Virtual planets
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## GCM : General Circulation Models

First GCMs were designed in late 50's, early 60's

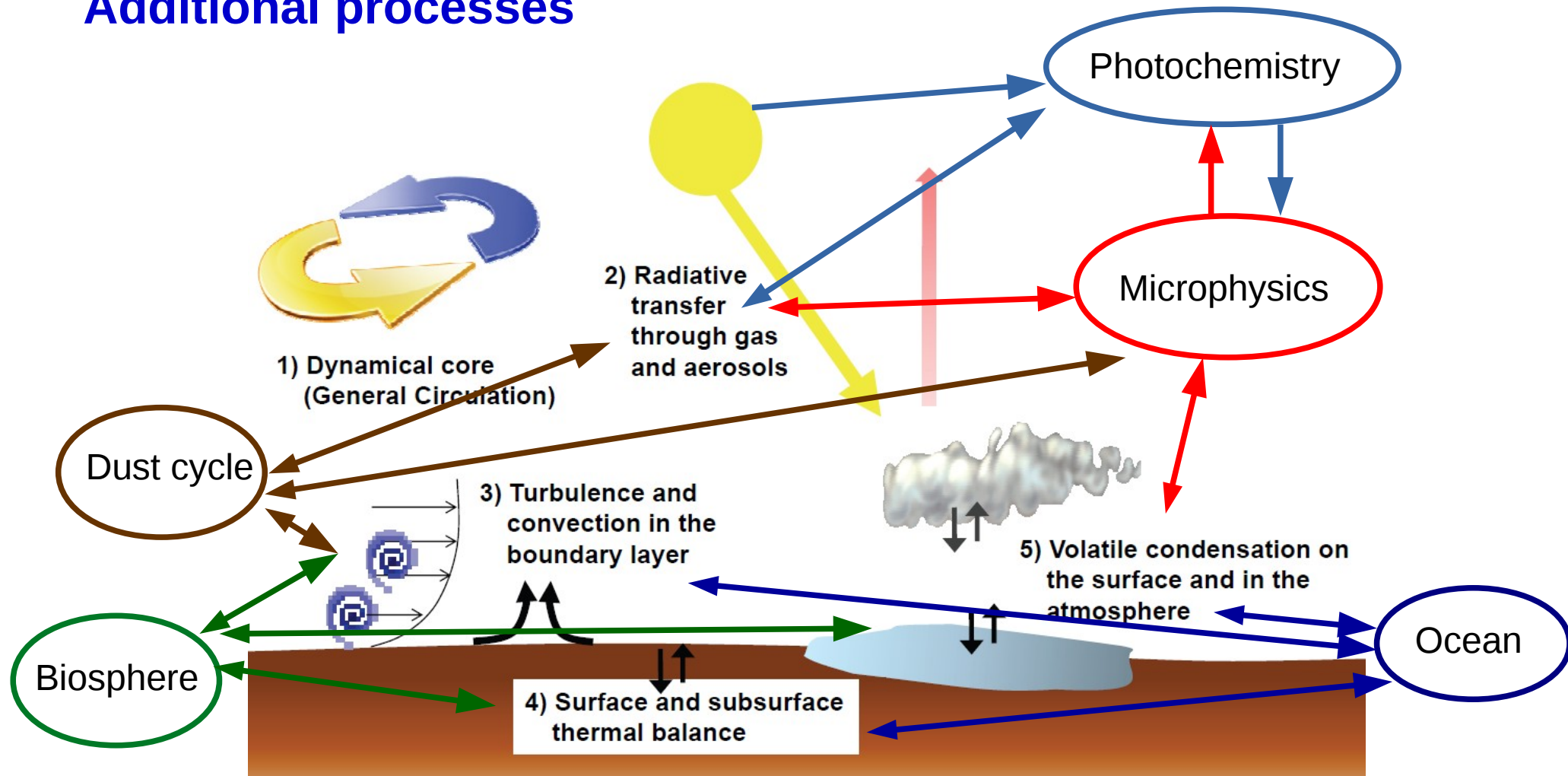
**Design : mandatory bricks**



## GCM : Global Climate Models

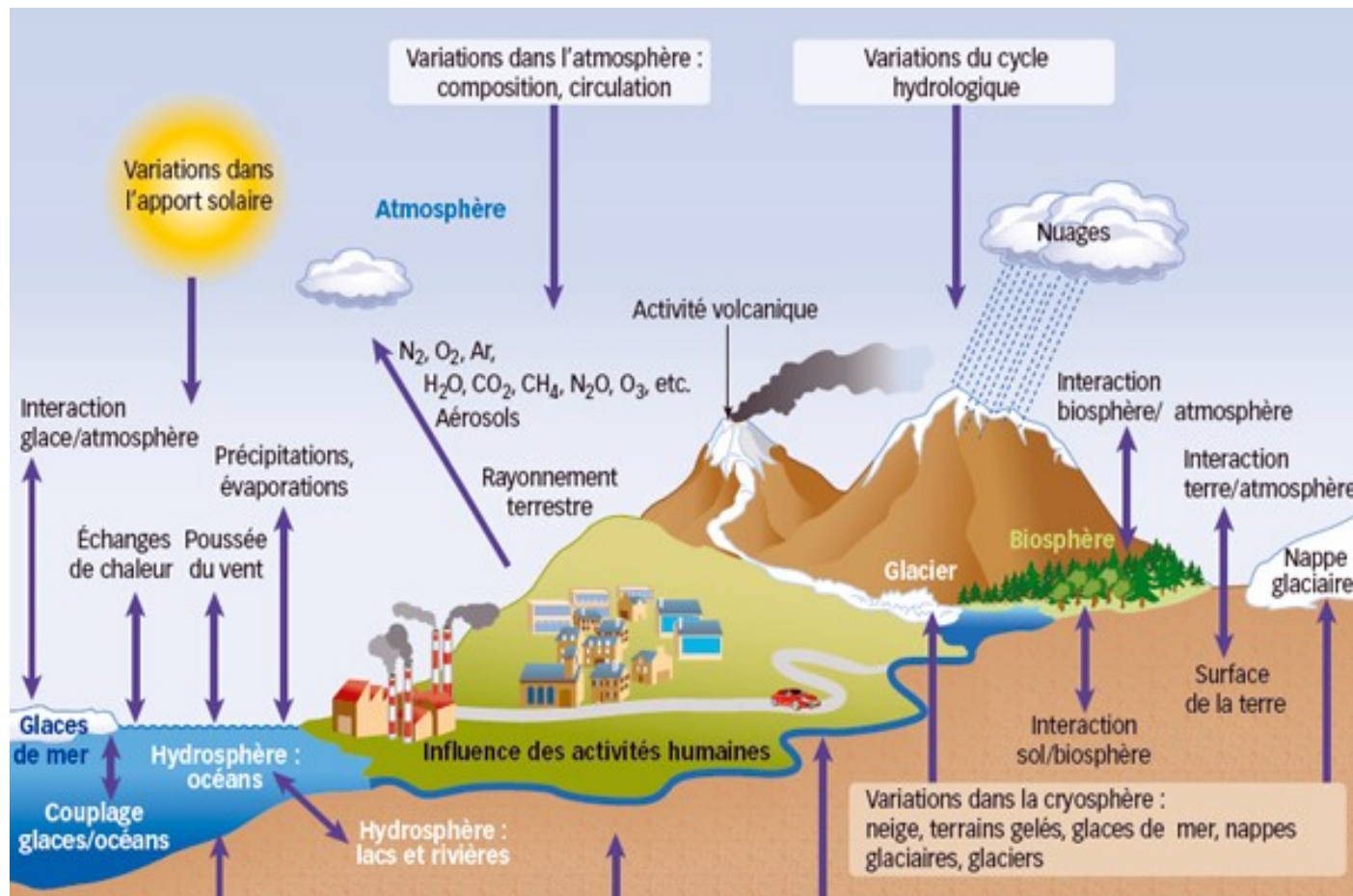
Because an atmosphere is a complex, coupled system

### Additional processes



## GCM : Global Climate Models

### Earth



## The dynamical core

6 equations

- 6 variables, **dynamical** and **thermodynamical**
- **Forcings** and **planetary constants**

Zonal momentum

$$\frac{du}{dt} - \frac{uv \tan \phi}{a} = 2\Omega \sin \phi v - \frac{1}{\rho} \frac{\partial p}{\partial x} + F_u$$

Meridional momentum

$$\frac{dv}{dt} + \frac{u^2 \tan \phi}{a} = -2\Omega \sin \phi u - \frac{1}{\rho} \frac{\partial p}{\partial y} + F_v$$

Hydrostatic balance

$$-\frac{1}{\rho} \frac{\partial p}{\partial z} - g = 0$$

Mass conservation

$$\frac{\partial \rho}{\partial t} + \text{div}(\rho \mathbf{V}) = 0$$

First principal of thermodynamics

$$\frac{c_p}{\theta} \frac{d\theta}{dt} = \frac{Q}{T} \quad \text{with} \quad \theta = T \left[ \frac{p_0}{p} \right]^\kappa$$

Ideal gas equation of state

$$p = \rho R T$$



## The dynamical core

### Different types of GCM

Hydrostatic vs Quasi-hydrostatic

Shallow atmosphere vs Deep atmosphere

$C_p(T)$ ,  $C_p(\text{composition})$

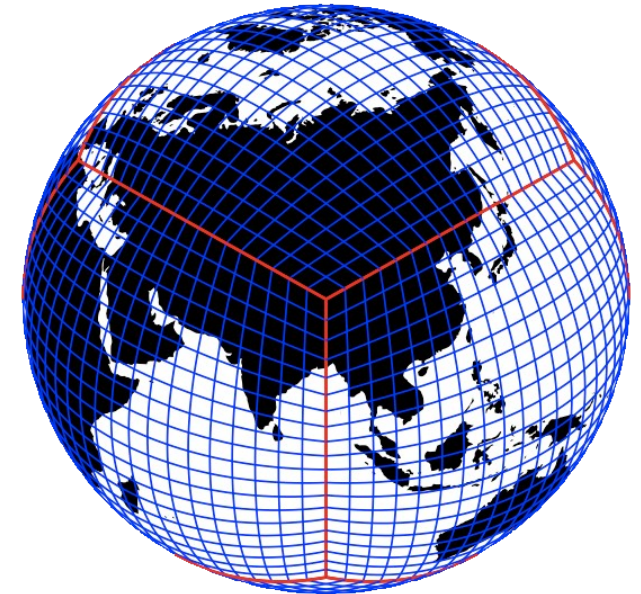
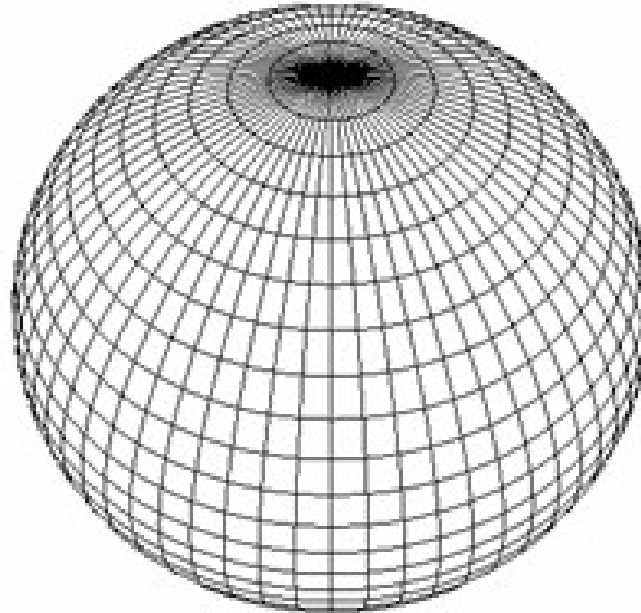
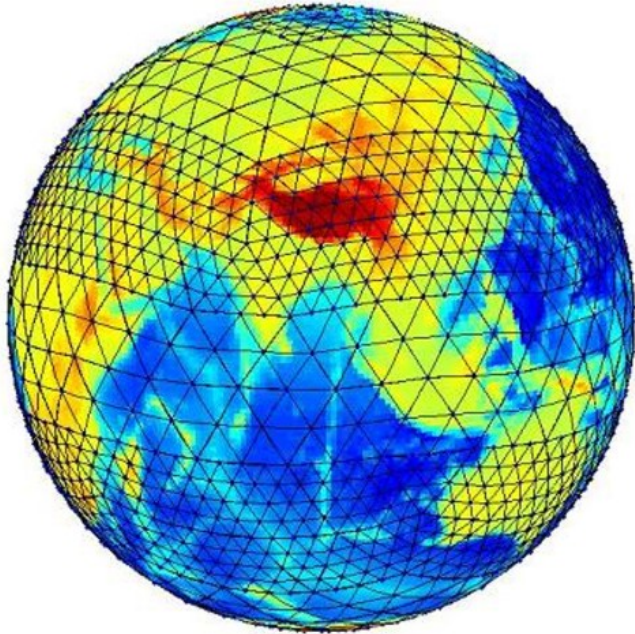


## The dynamical core

### Horizontal discretization

Finite differences, finite volumes

Spectral (spherical harmonics), spectral elements



Vertical discretization : mass, altitude, pressure...

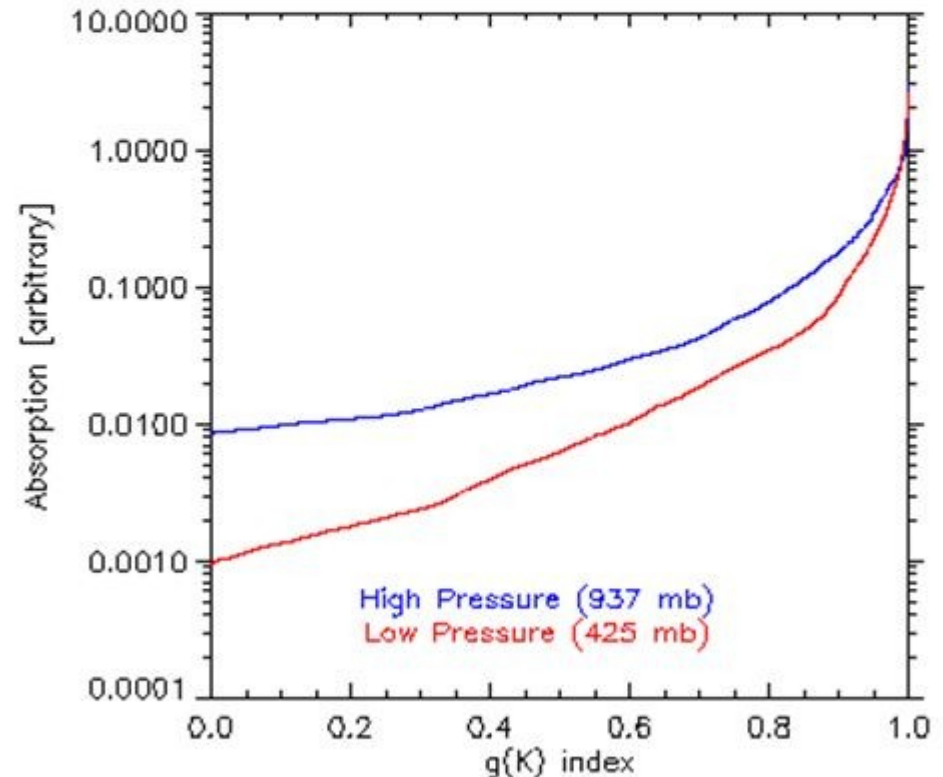
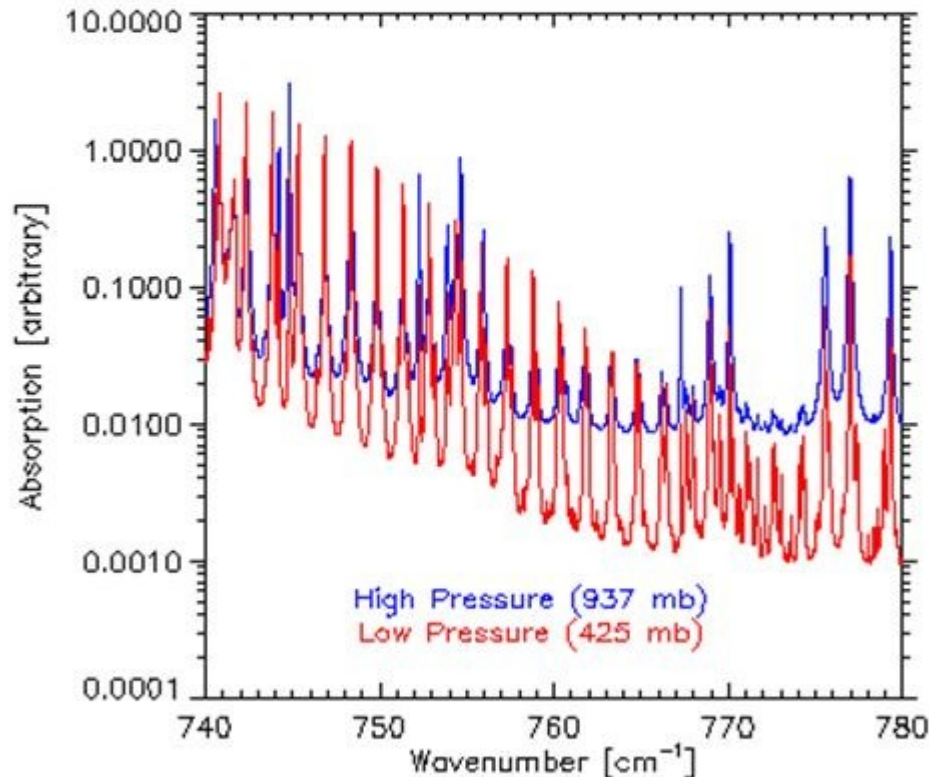
Time marching scheme

Conservation concerns : mass, energy, angular momentum

## The radiative transfer

**Strategy** : fast computation ; versatility

Usual approach to gas opacities : correlated-k distribution



## Physics processes

Table 1: Summary of key processes and related problems in low atmosphere terrestrial planet global climate models.

Key physical processes	Earth	Mars	Venus	Titan	Triton/Pluto
Radiative transfer	×	×	Optical thickness <sup>a</sup>	×	× <sup>b</sup>
Clouds	H <sub>2</sub> O	H <sub>2</sub> O ice, CO <sub>2</sub> ice	H <sub>2</sub> SO <sub>4</sub>	CH <sub>4</sub> , C <sub>2</sub> H <sub>6</sub>	N <sub>2</sub> ice?
Hazes	Aerosols	Mineral dust		Organic haze	Organic haze?
Turbulence and convection	Near-surface	Near-surface	Cloud layers	Near-surface	Near surface
Subsurface heat storage	With oceans	×	×	×	Long-term buffer <sup>c</sup>
Dominant gas condensation		CO <sub>2</sub>			N <sub>2</sub>
Minor species condensation	H <sub>2</sub> O	H <sub>2</sub> O		CH <sub>4</sub> , C <sub>2</sub> H <sub>6</sub>	CH <sub>4</sub> ?
Dynamical core					
Deep atmosphere				×	
Specific heat variations			$C_p(T)^e$		
Composition variations		×			
Momentum conservation			Critical <sup>d</sup>	Critical <sup>d</sup>	

*a*: Eymet et al. (2009) ; *b*: Triton's atmosphere N<sub>2</sub> atmosphere is so pure and tenuous that gaseous absorption of radiation can be neglected; Molecular conduction then dominates heat transport. (Yelle et al. 1991; Vangichith et al. 2013). *c*: see section 2.5 ; *d*: see section 2.2 ; *e*: Lebonnois et al. (2010a).

# PLANETARY ATMOSPHERES

## Global Climate Modeling

- Virtual planets
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- Successes, and lessons from failure

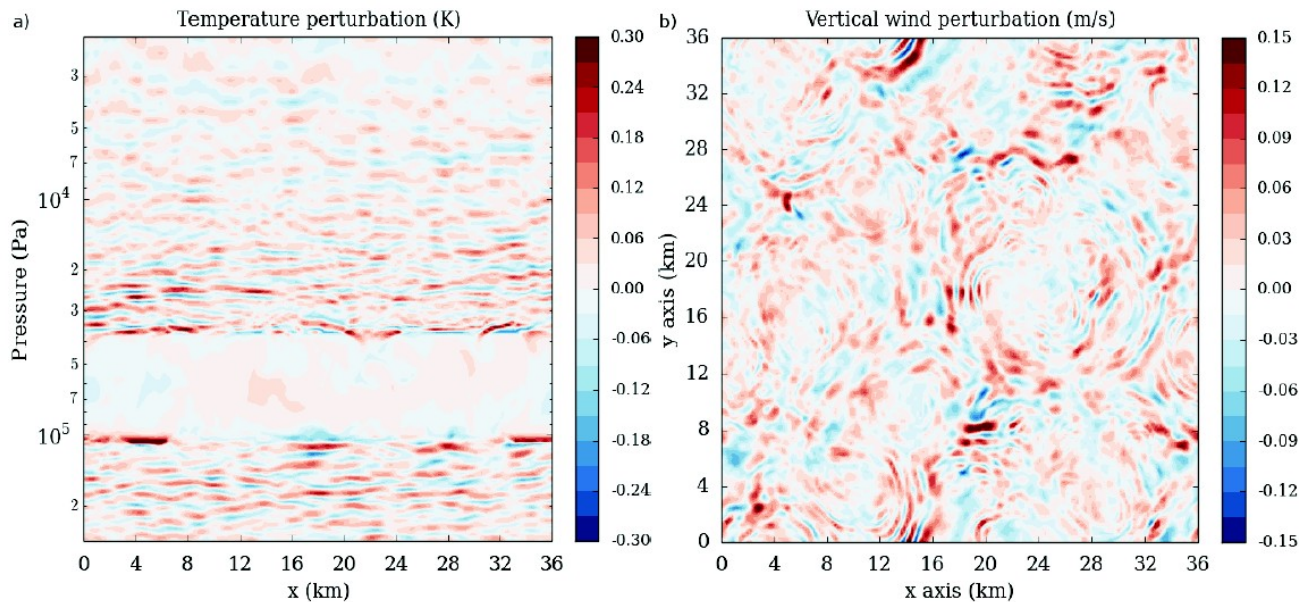


## Small-scale processes

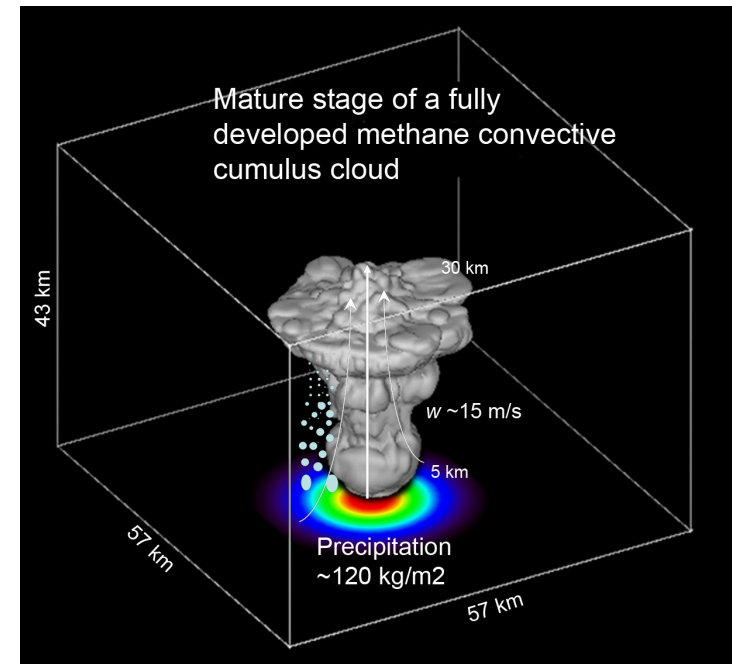
### Large Eddy Simulations :

small scale (1-1000m), idealized, non-hydrostatic models

=> study small-scale processes :  
turbulence, convection  
gravity waves



Venus convective layer in cloud : gravity waves

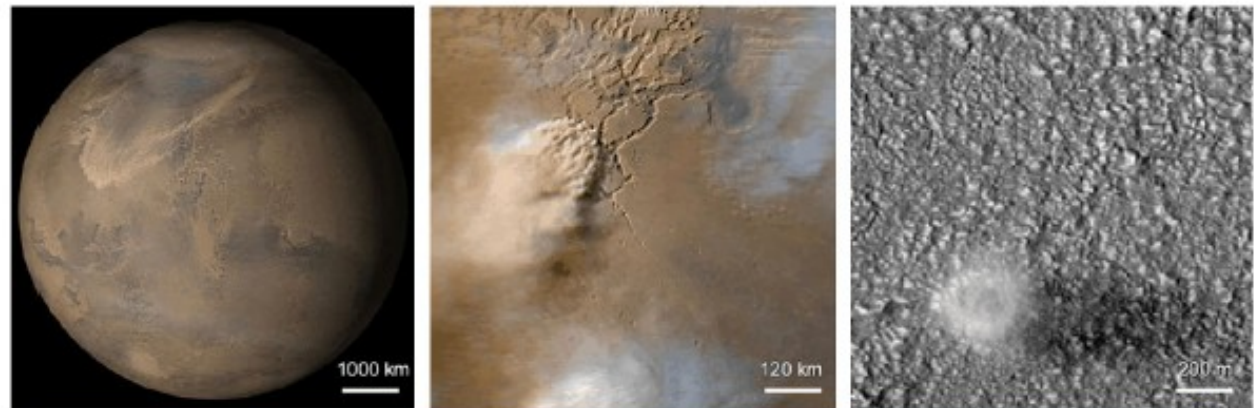


Titan methane convective clouds

## Intermediate-scale processes

### Regional Climate Models :

- Intermediate scales => small-scales parameterized
- non-hydrostatic
- boundary conditions from GCMs



... Dust fronts ... Regional dust storms ... Local gusts ... Dust devils ...

10000 km 1000 km 100 km 10 km 1 km 100 m 10m 1m

Global Circulation Models

Mesoscale Models

Large-Eddy Simulations

## Intermediate-scale processes

### Regional Climate Models

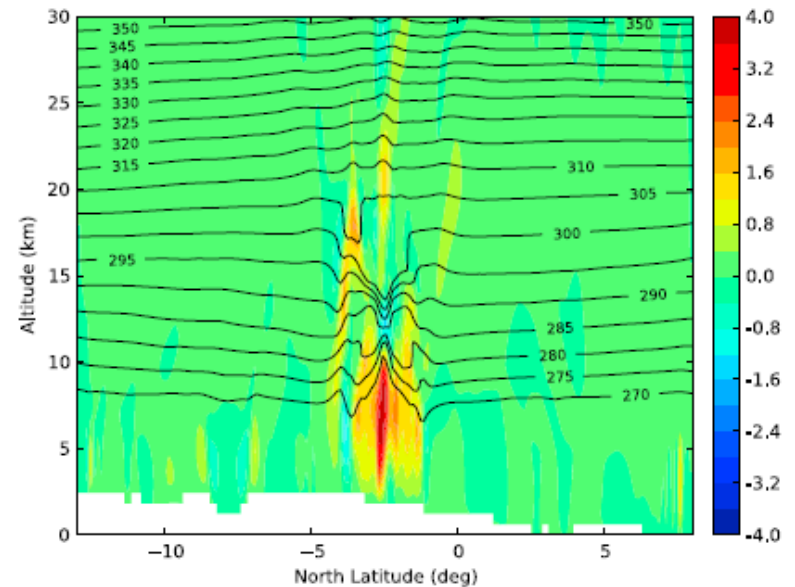
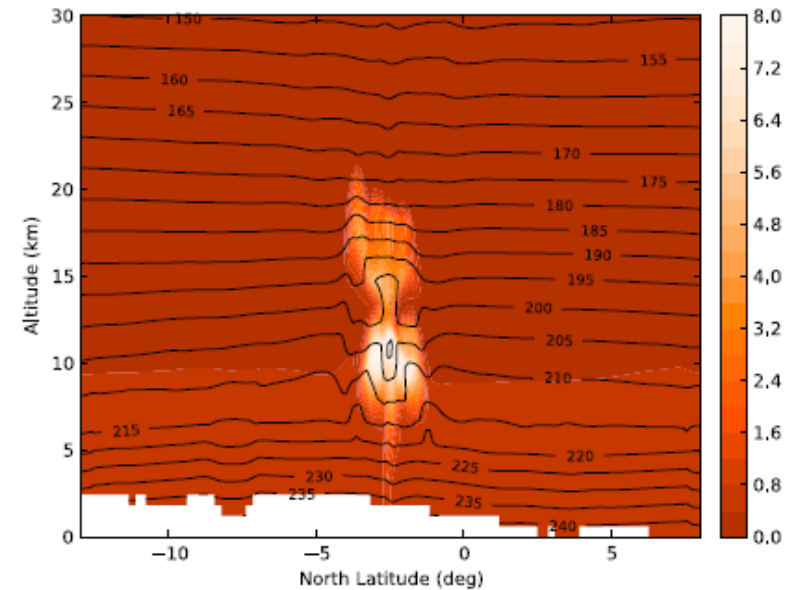
Dust optical depth

Temperature

Example : rocket dust storm on Mars  
(Spiga et al, 2013)

Vertical wind

Potential temperature





# PLANETARY ATMOSPHERES

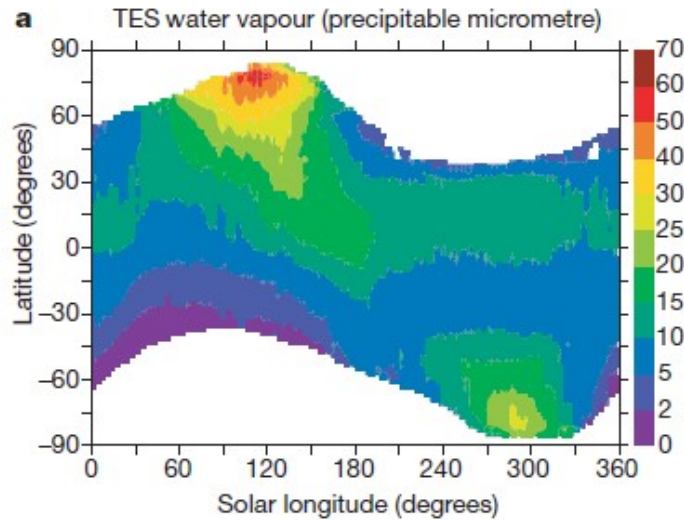
## Global Climate Modeling

- Virtual planets
- Different models for different scales
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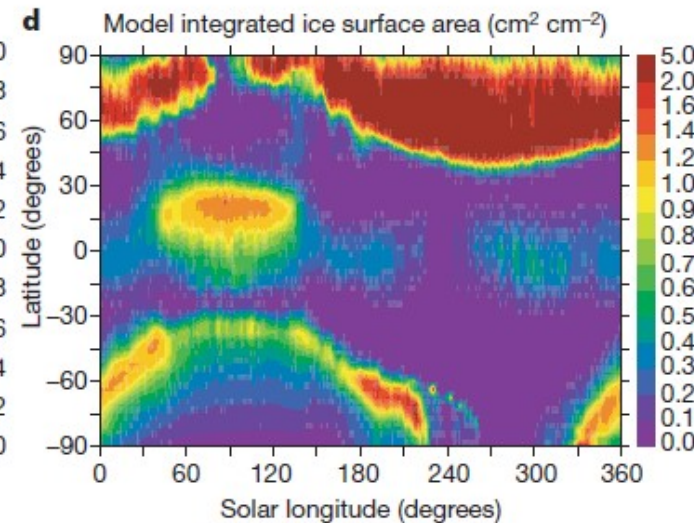
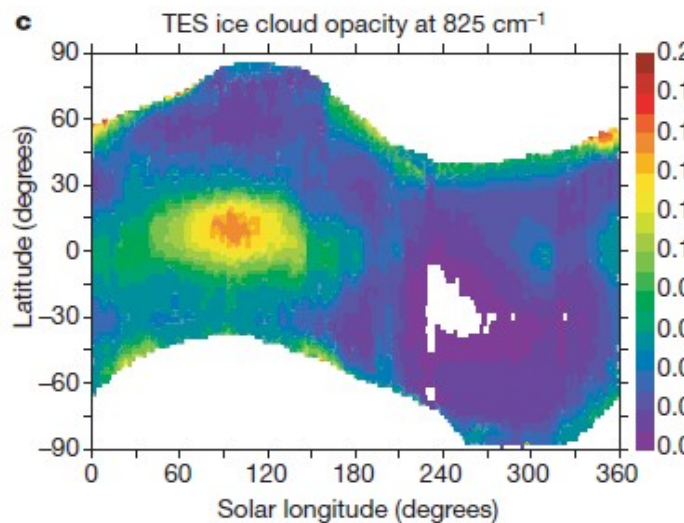
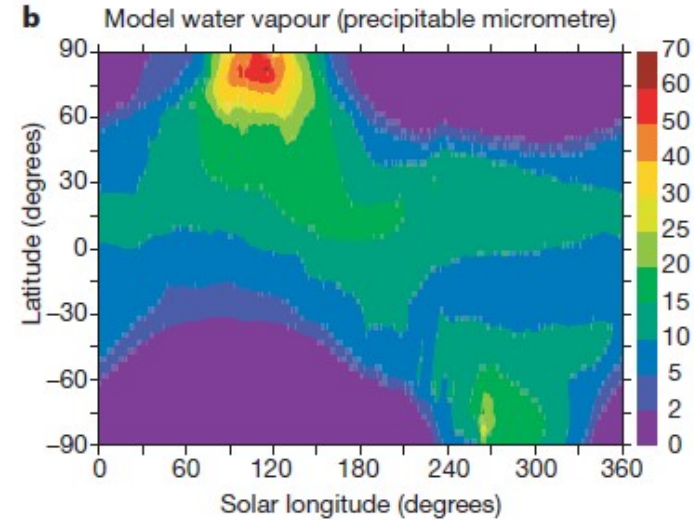
## Mars climate

### Water and cloud cycles

TES observations



GCM simulations



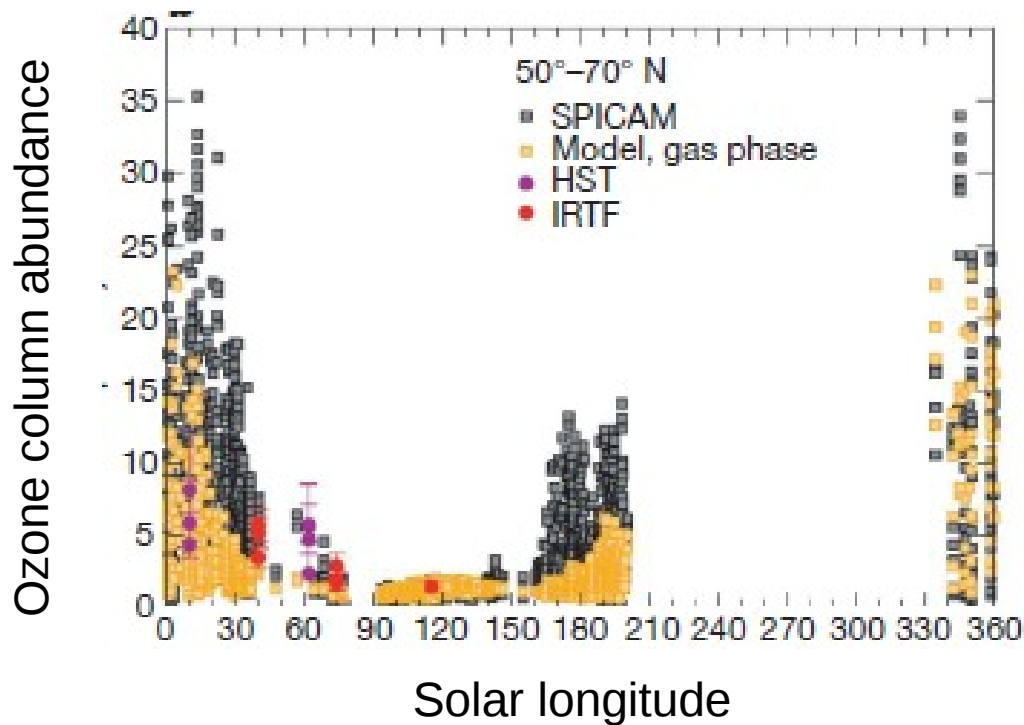
Water vapor

Clouds

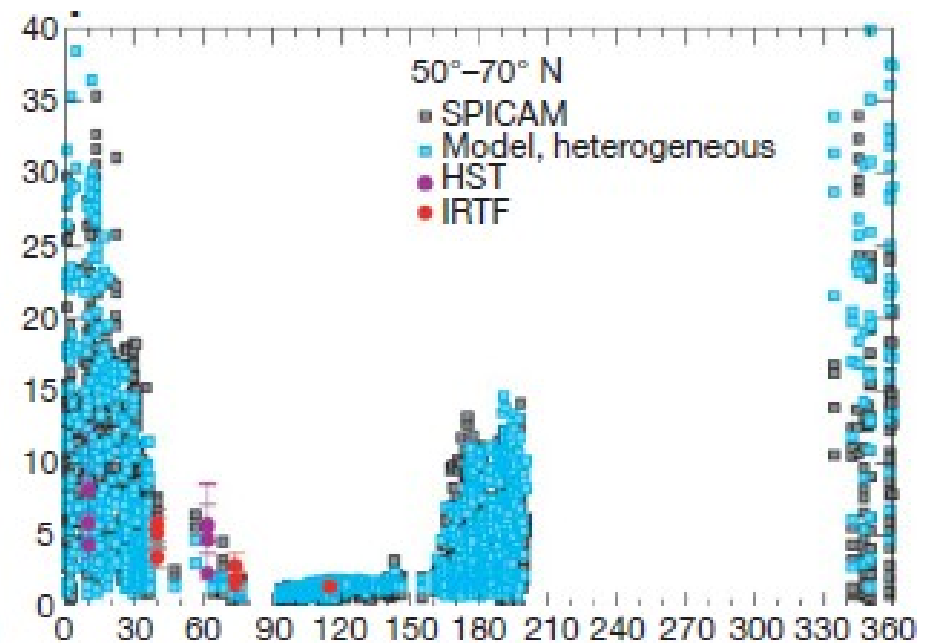
## Mars climate

### Heterogeneous chemistry

Only gas photochemistry



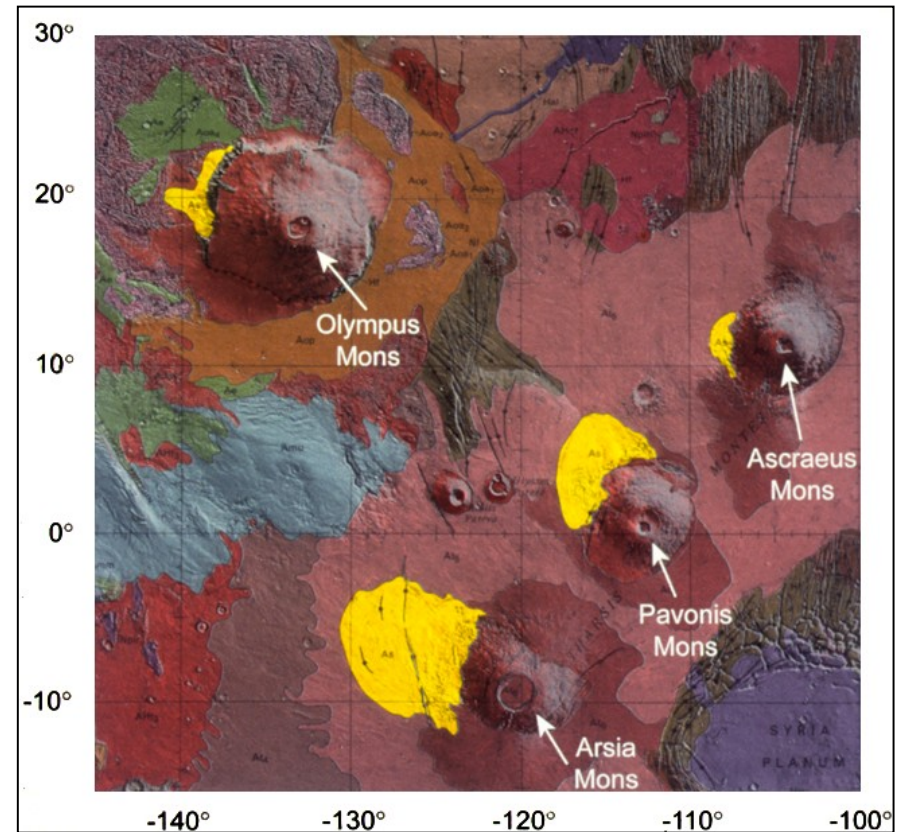
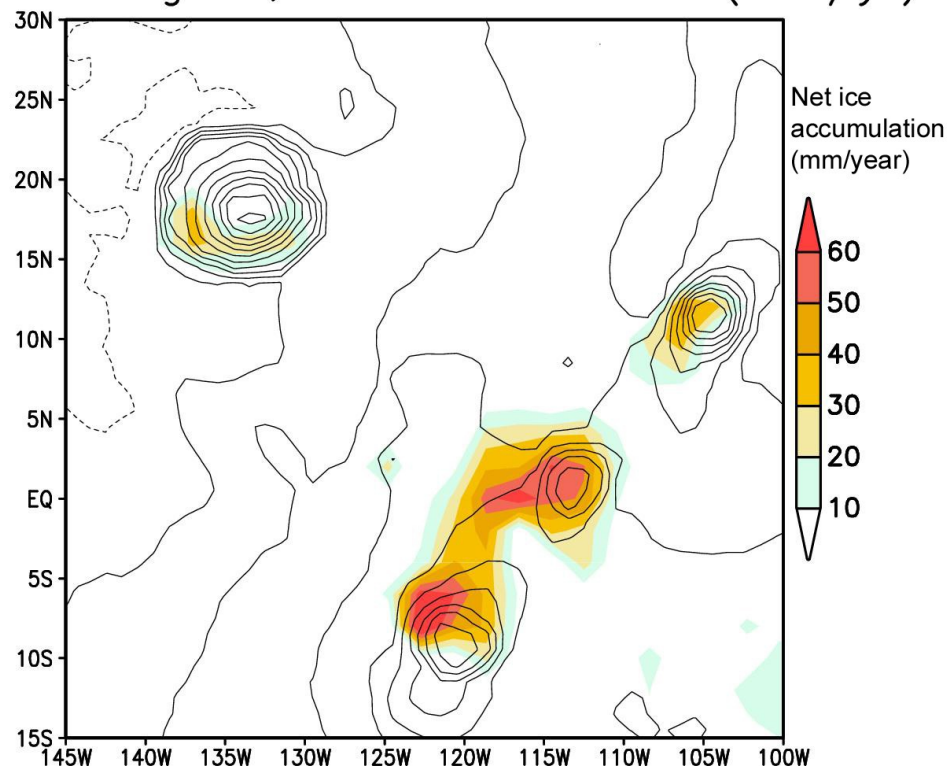
With heterogeneous chemistry on ice cloud particles



## Mars paleoclimate

### Non-polar glaciers, millions of years ago

Glacier formation :  
Ice accumulation rate (mm/y)  
with obliquity=45°

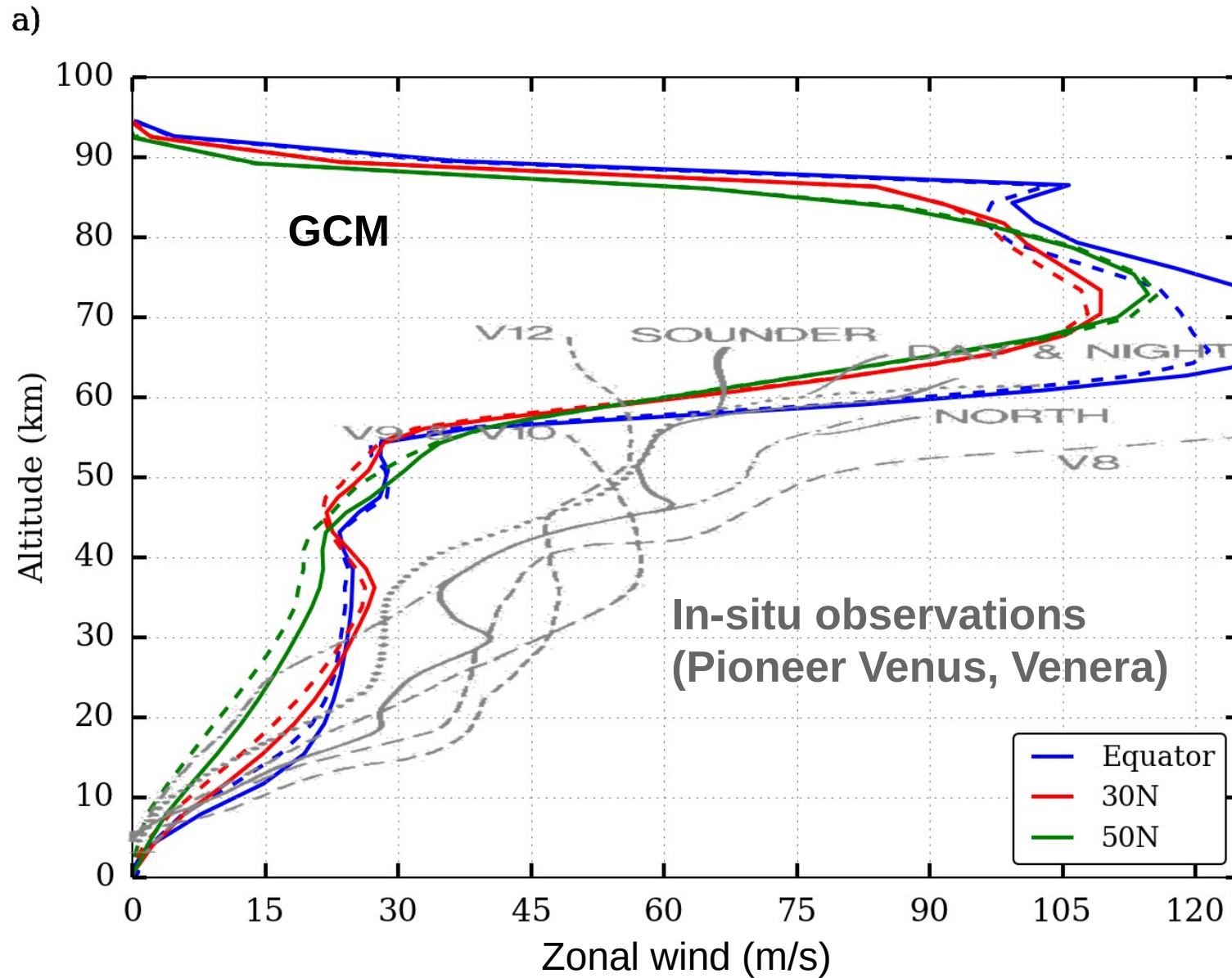


Fan shaped deposits, drop moraines characteristic of cold based glaciers.

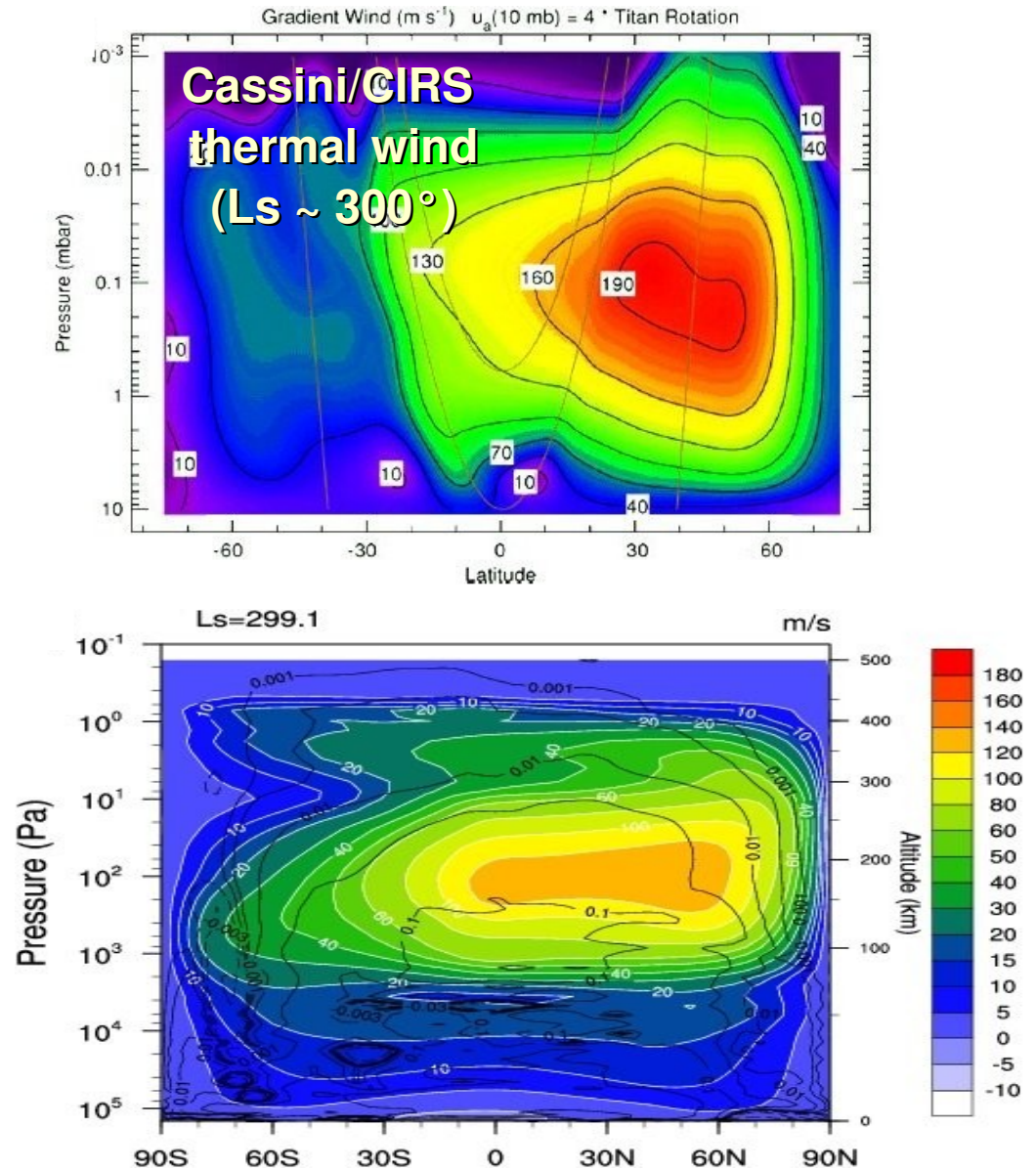
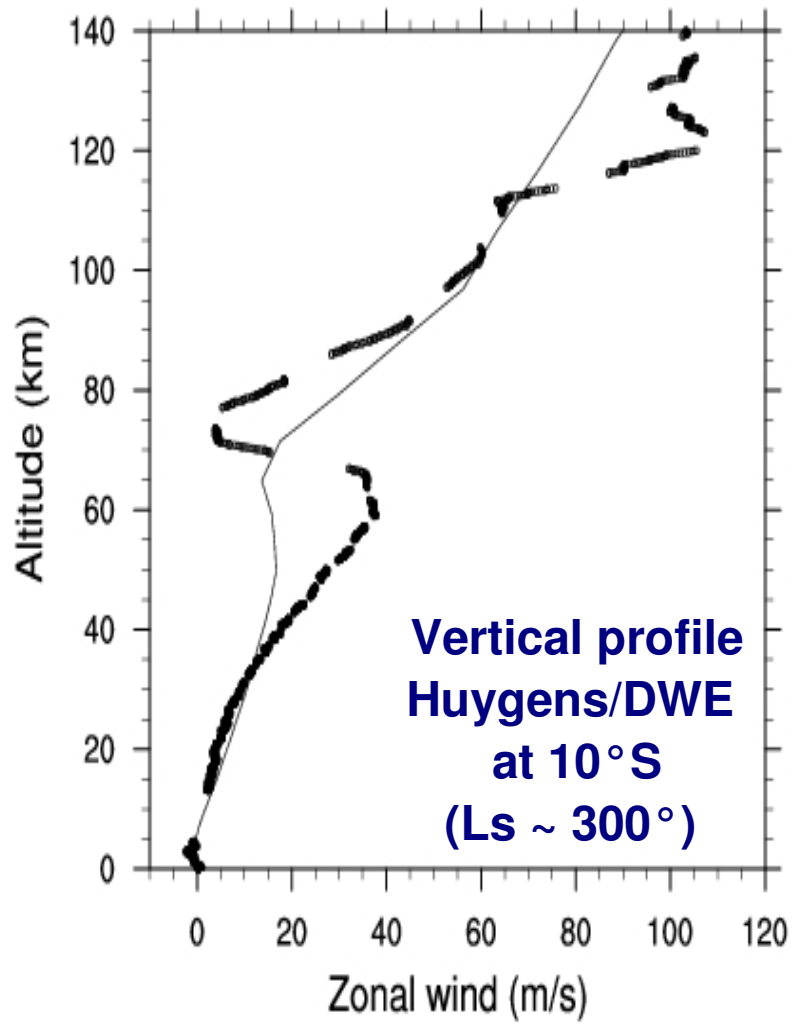
Rock glaciers



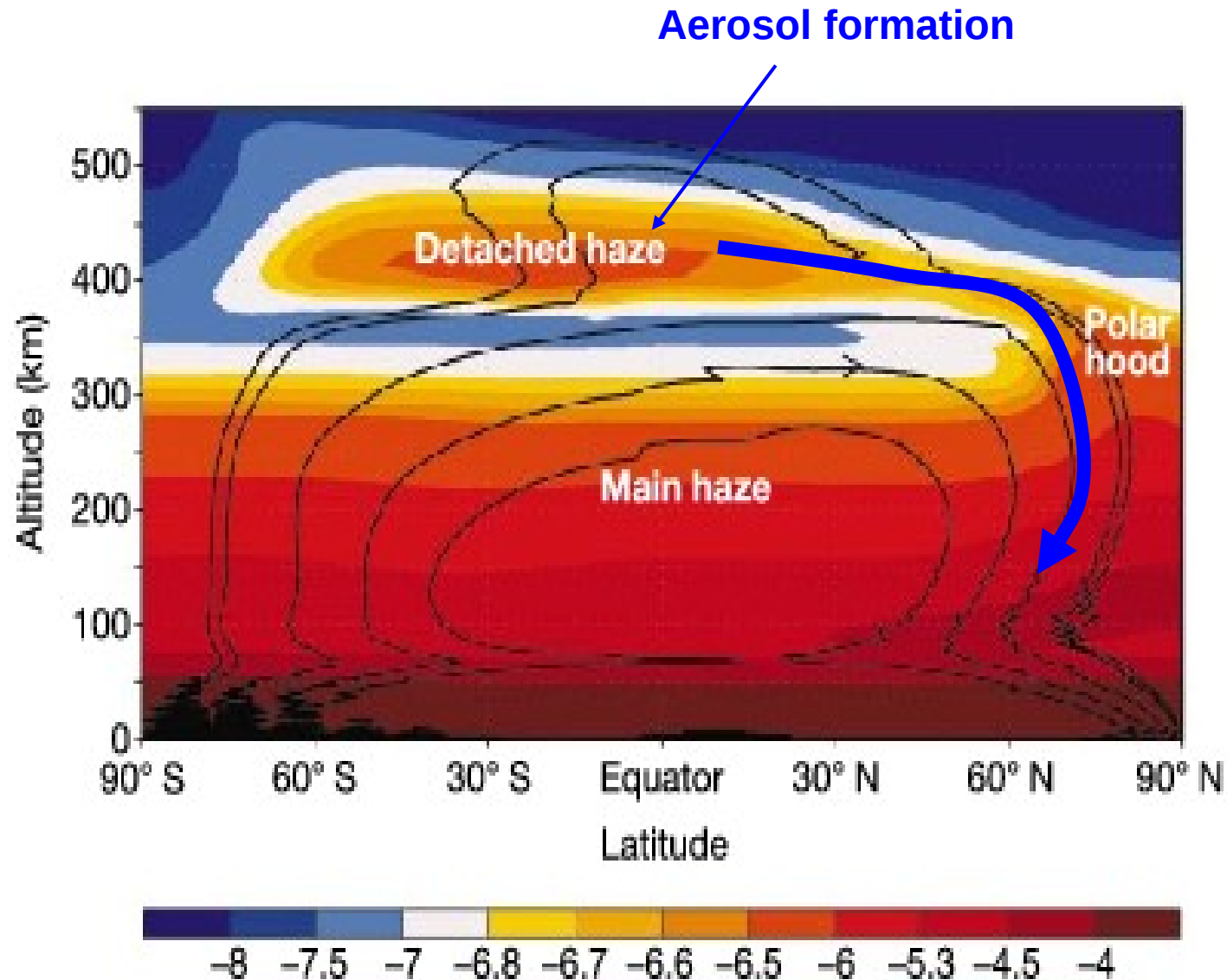
## Venus superrotation



## Titan superrotation

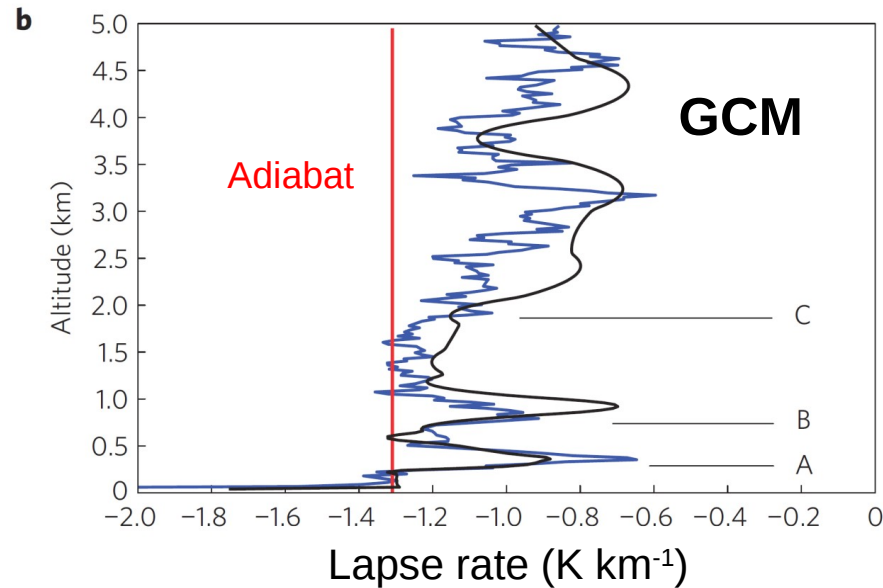
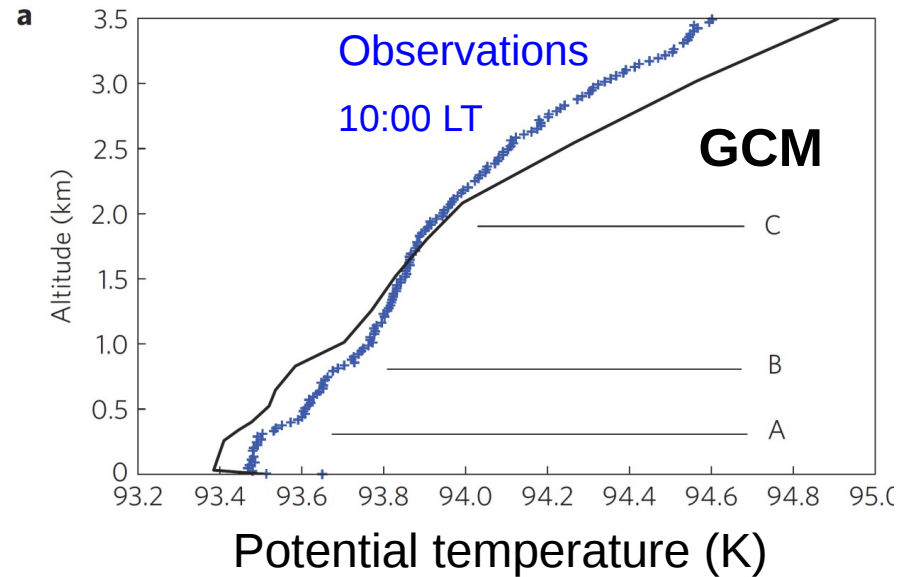


## Titan detached haze layer



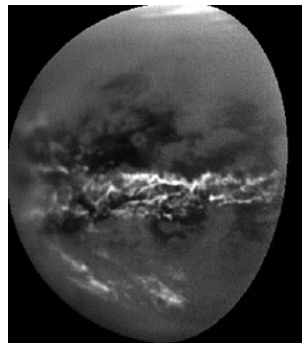
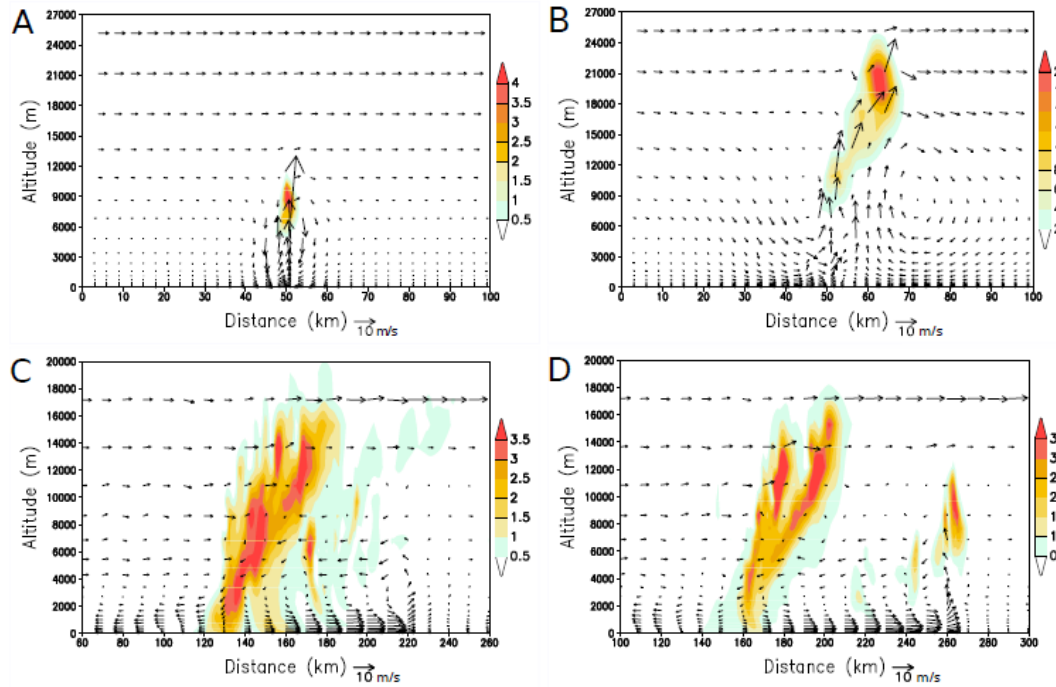


## Titan boundary layer

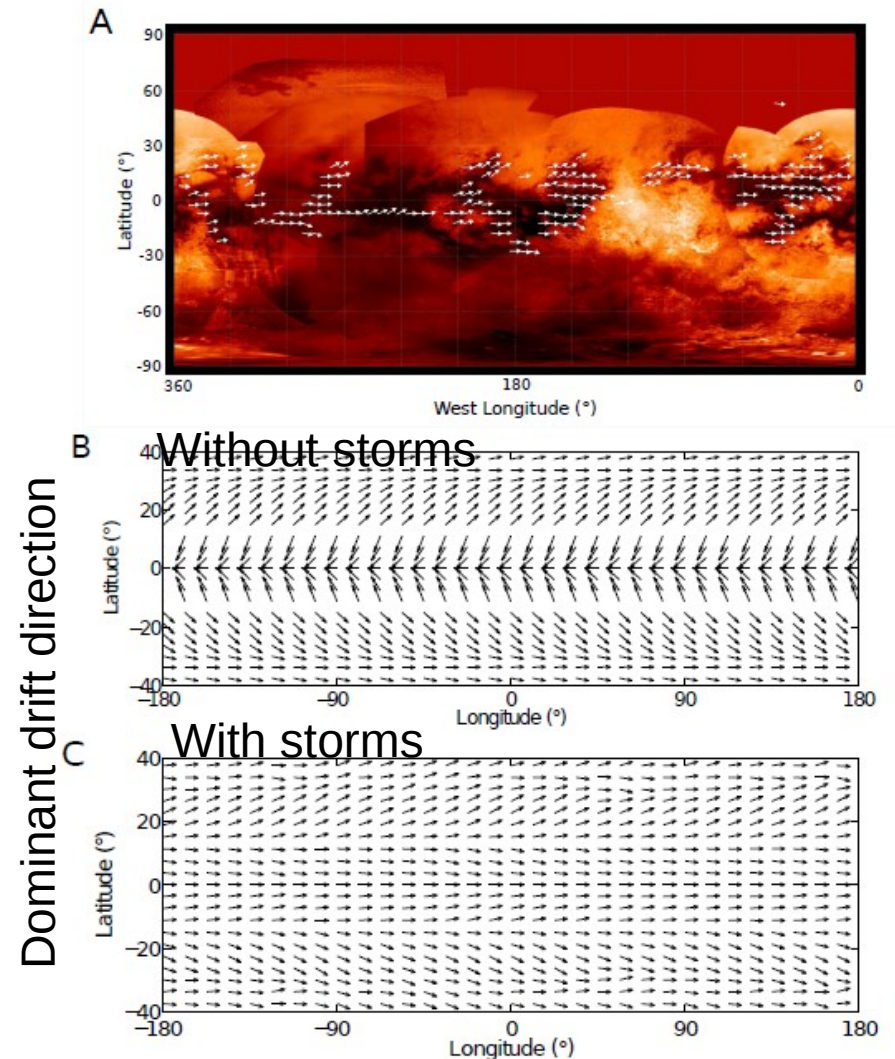


## Titan equatorial dunes

Mesoscale (2D only) simulation of an equinox storm



Dunes orientation



### Lessons from failure

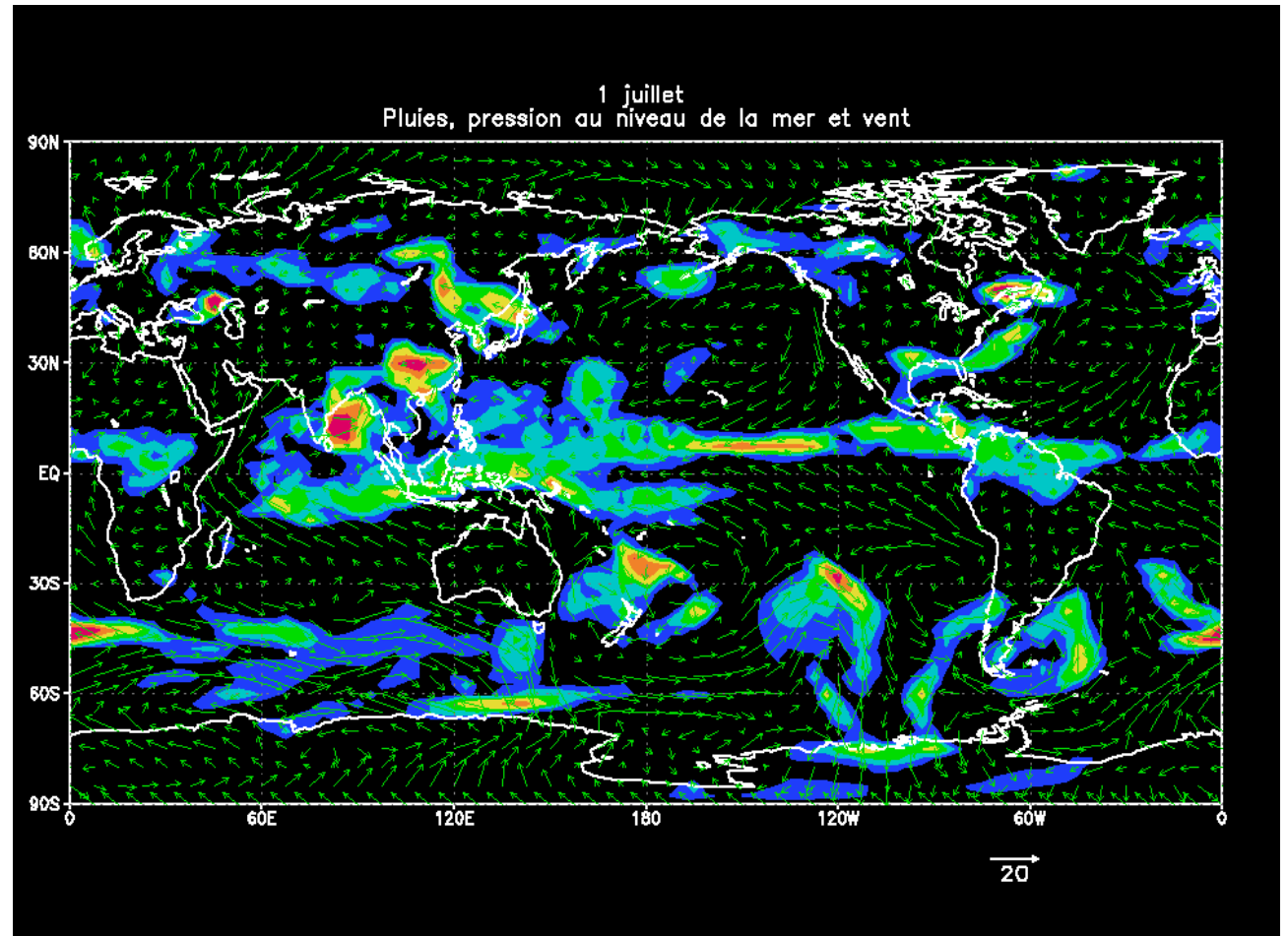
- Missing physical processes **Radiative effect of Martian clouds**
- Insufficient representation of physical processes :
  - complex sub-scale processes **Terrestrial clouds – gravity waves**
  - Positive feedbacks, instabilities **Rétroaction due to ice albedo**
  - Non-linearities, thresholds **Martian dust storms**
- Long time scales, sensitivity to initial state **Pluto ices**
- Weak forcings : when the system evolution is sensitive to a subtle balance between processes, and not driven by a strong forcing

### **Superrotation**

## Terrestrial clouds

Complex microphysics  
and small-scale dynamics

Precipitations in a GCM  
Global scale, coarse resolution (~100 km)

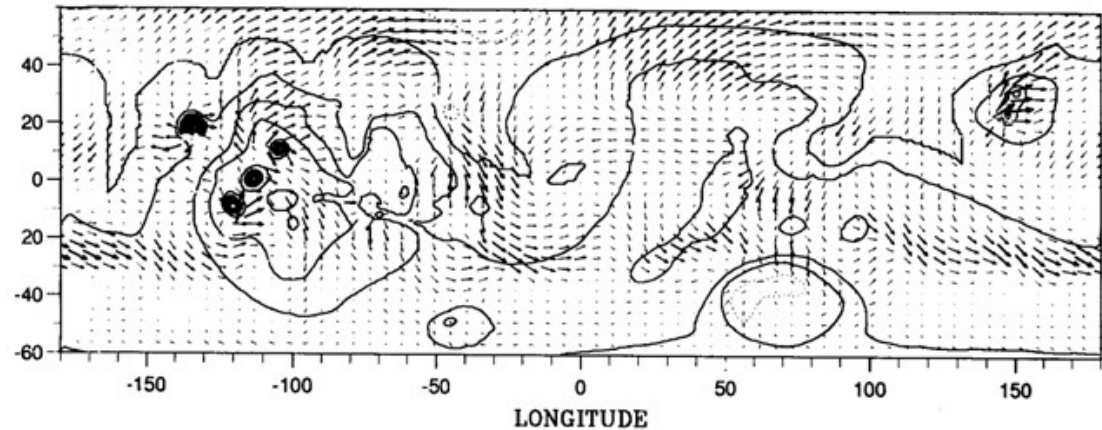




## Martian dust storms

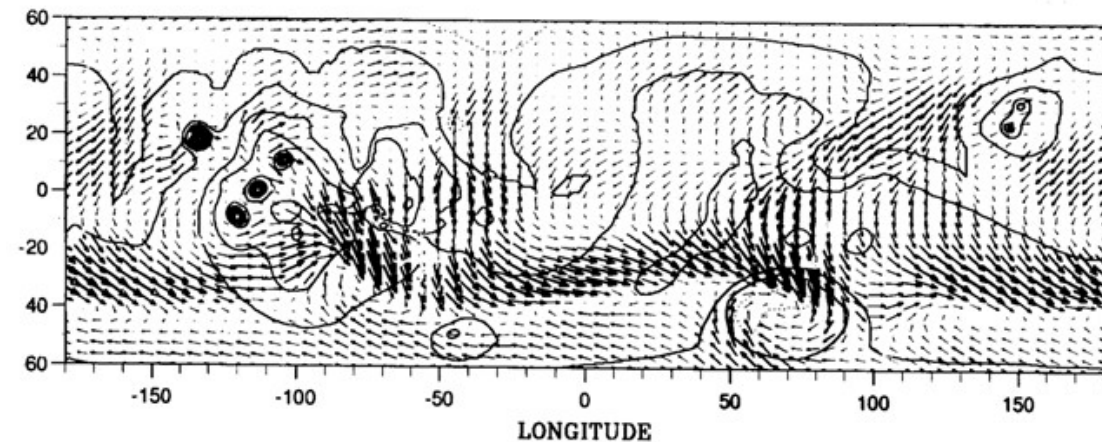
Strong positive retroaction of dust on circulation and on dust lifting

Clear atmosphere



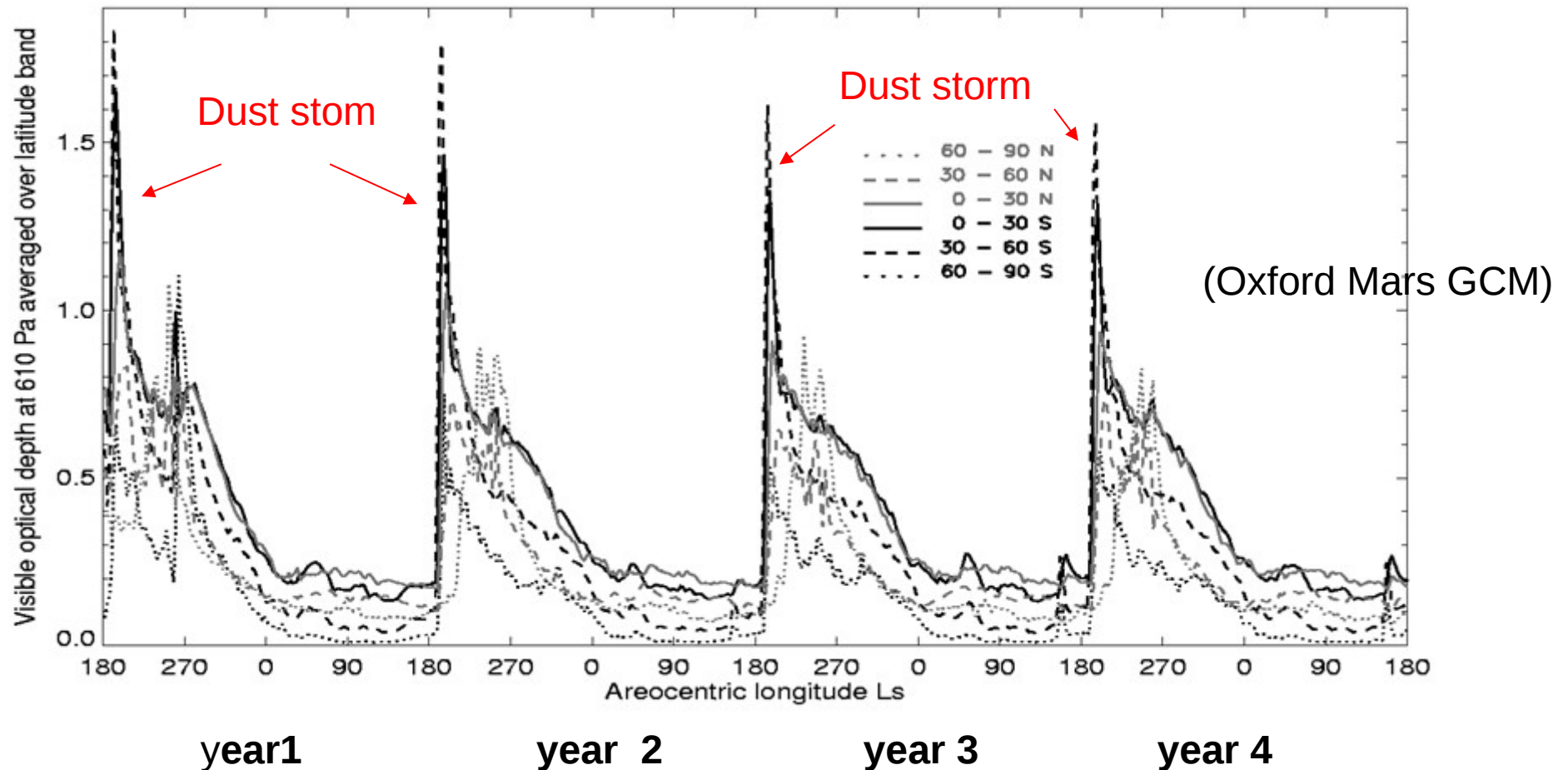
(LMD Mars GCM)

Dust storm



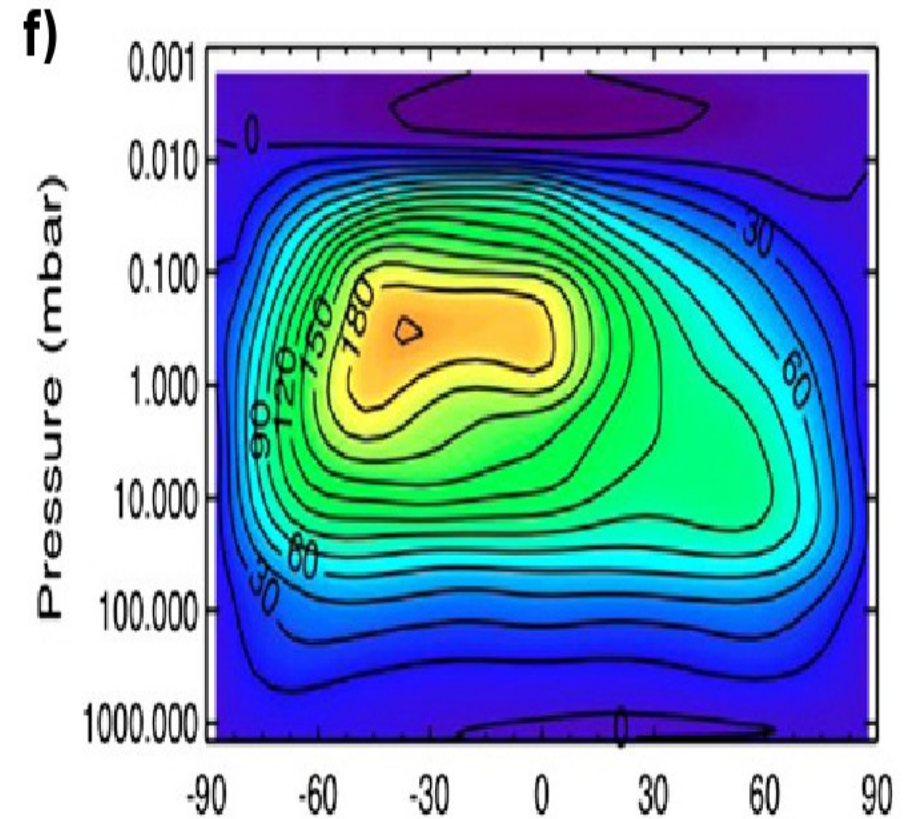
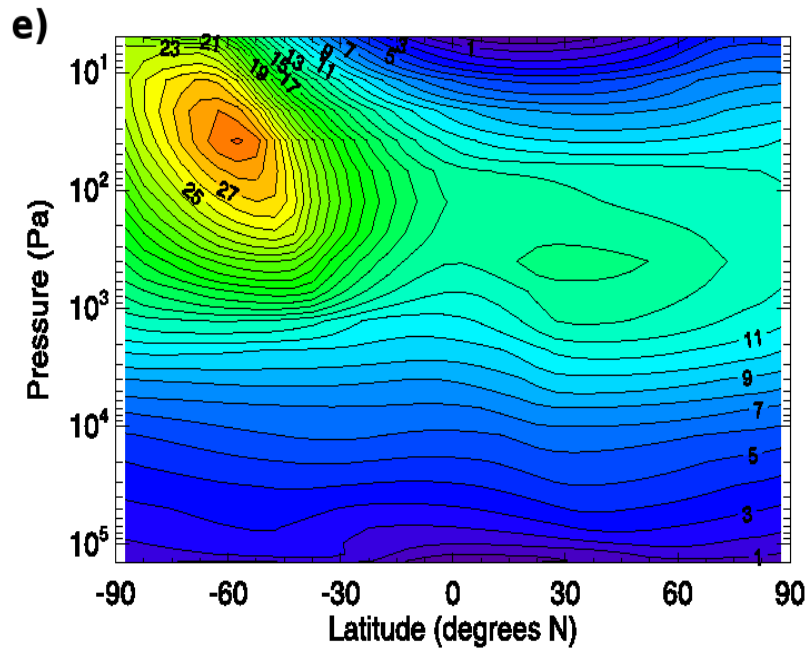
## Martian dust storms

Difficult to reproduce the interannual variability...



## Titan superrotation

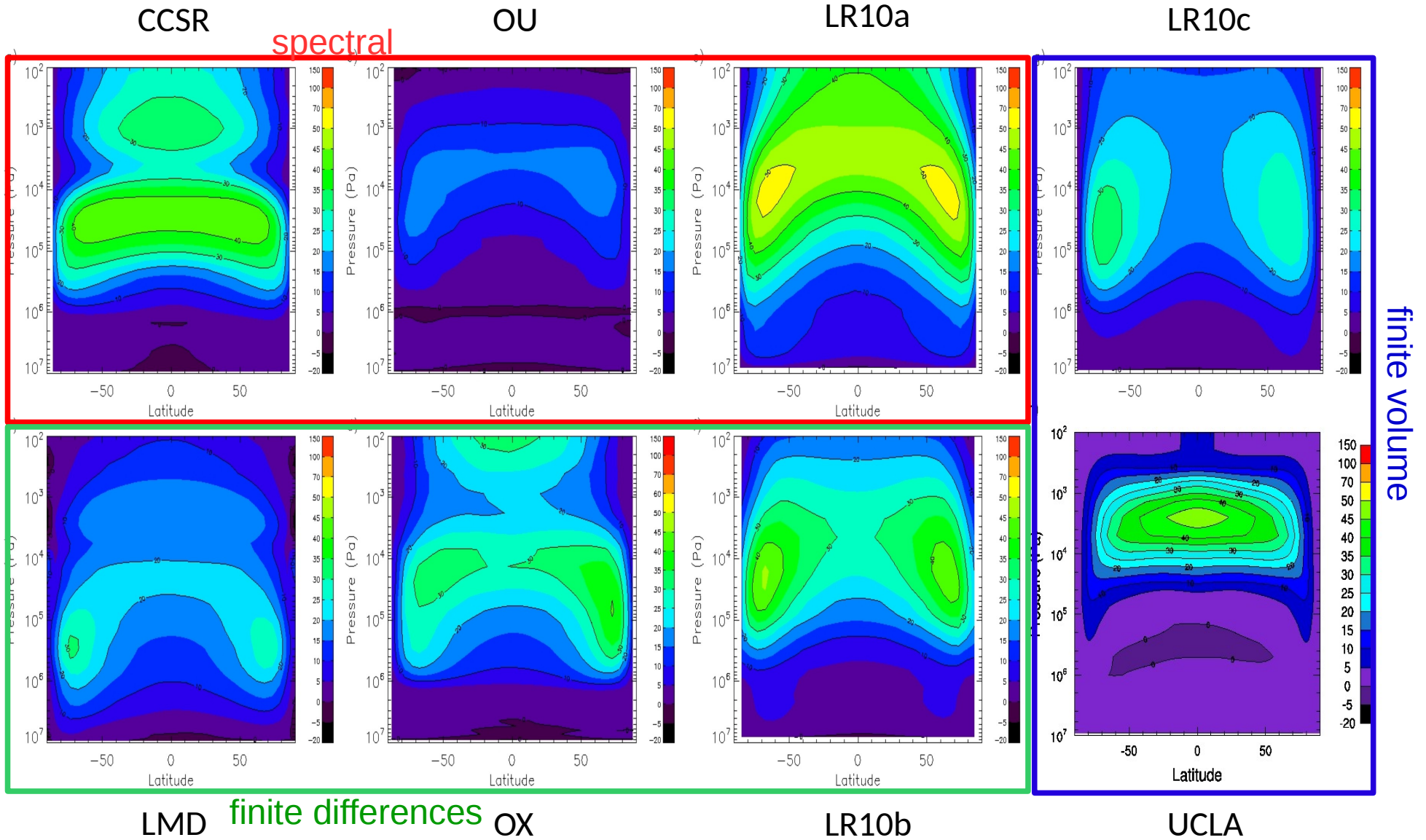
Same GCM, corrections in the dissipation formulation





## Venus superrotation

Several dynamical cores, same simplified physics



## To conclude

### Suggested bibliography

- A. Sánchez-Lavega, *An Introduction to Planetary Atmospheres*, CRC Press, Taylor and Francis, 2011, ISBN 9781420067354.
- Forget F. and Lebonnois S., « Global climate models of the terrestrial planets »  
+ Dowling T., « Earth General Circulation Models »  
in *Comparative Climatology of the Terrestrial Planets*, S.J. Maxwell et al. Eds.  
University of Arizona, Tuscon, 2013

### Acknowledgements

- thanks to the authors of all the images and plots I showed.  
I can send references upon request
- thanks to Julia and Javier for the invitation and this great Winter School.