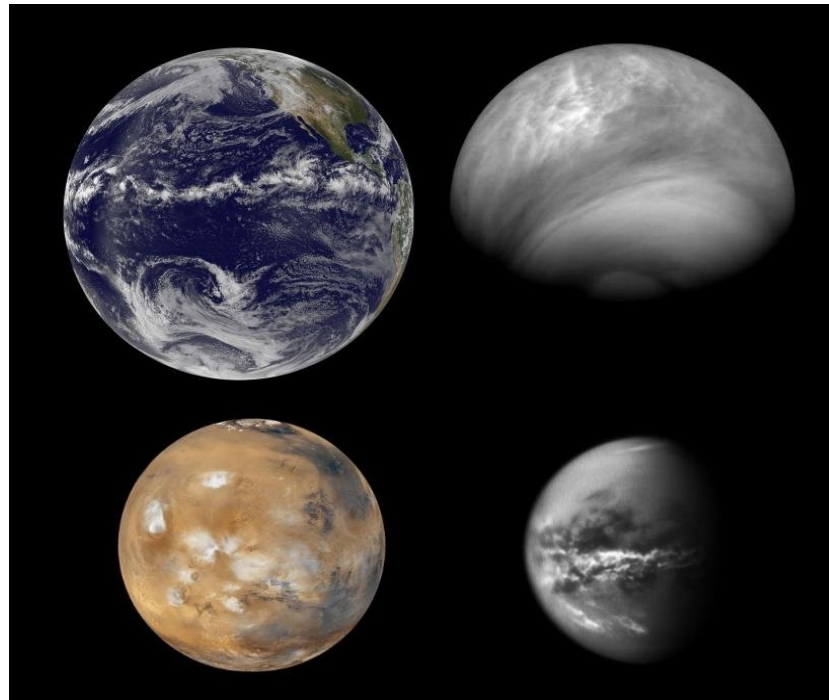


# PLANETARY ATMOSPHERES

## 3. Atmospheric dynamics and circulation regimes



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## PLANETARY ATMOSPHERES

### Atmospheric dynamics and circulation regimes

- Equations of the atmospheric fluid
- Circulation patterns in terrestrial atmospheres
- Instabilities and wave activity
- Vortices

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## Basic laws

### Equation of state : ideal gas law

We consider the atmospheric gas as **ideal** :

$$p = \rho R T$$

pressure                      density                      temperature

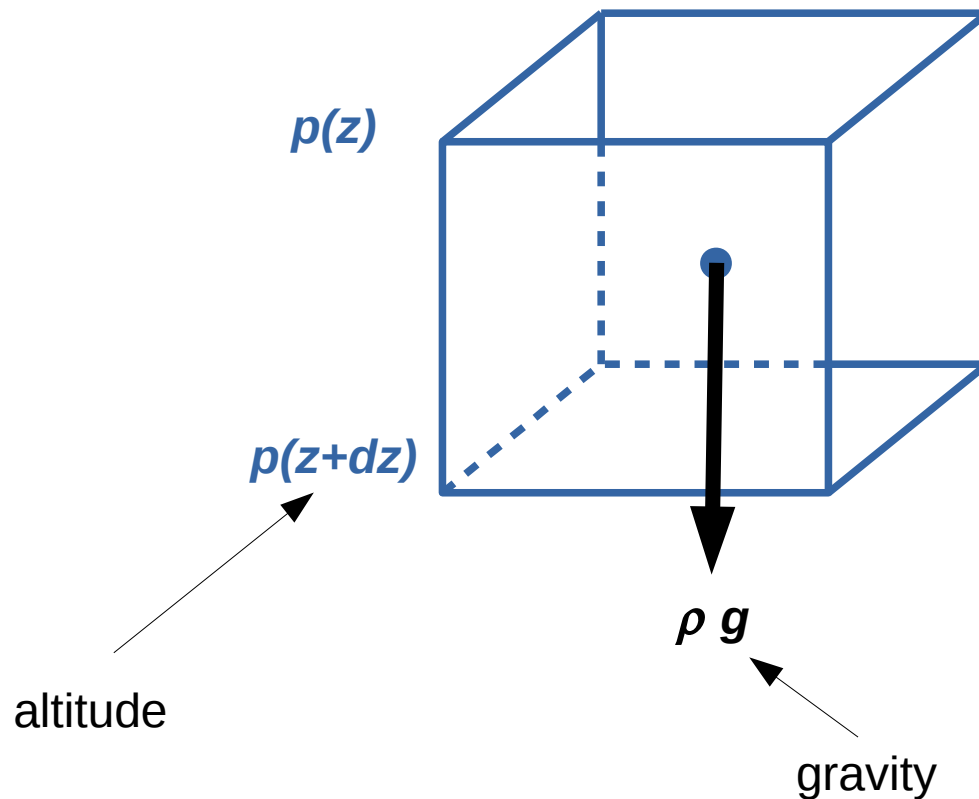
mean molecular mass

$$R = R^* / \mu$$
$$R^* = 8.314 \text{ J mol}^{-1} \text{ K}^{-1}$$

## Basic laws

### Hydrostatic equilibrium

Fluid elemental particle **at rest**



$$\rho g + \frac{\partial p}{\partial z} = 0$$

## Basic laws

### Scale height

Combining ideal gas law and hydrostatic equilibrium :

$$\frac{dp}{p} = -\frac{g}{RT} dz \quad H = \frac{RT}{g}$$

When  $T$  is taken as constant (**isothermal** atmosphere) :

$$p = p_0 e^{-z/H}$$

## Basic laws

### Adiabatic lapse rate

For a parcel receiving the heat quantity  $\delta q$ , from first principle of thermodynamics we get :

$$c_p dT = \frac{dp}{\rho} + \delta q$$

Specific heat capacity at constant pressure

When this parcel moves **adiabatically** :

$$\frac{dT}{T} = \frac{R}{c_p} \frac{dp}{p} \quad \Rightarrow \quad \Gamma_d = \frac{dT}{dz} = -\frac{g}{c_p}$$

## Basic laws

### Potential temperature

Moving adiabatically a parcel from  $(p, T)$  to a reference pressure  $p_0$ , its temperature will be  $\theta$ , defined as the potential temperature.

To get the expression for  $\theta$ , we integrate this expression

$$\frac{dT}{T} = \frac{R}{c_p} \frac{dp}{p}$$

If  $c_p$  does not depend on  $T$ , we get

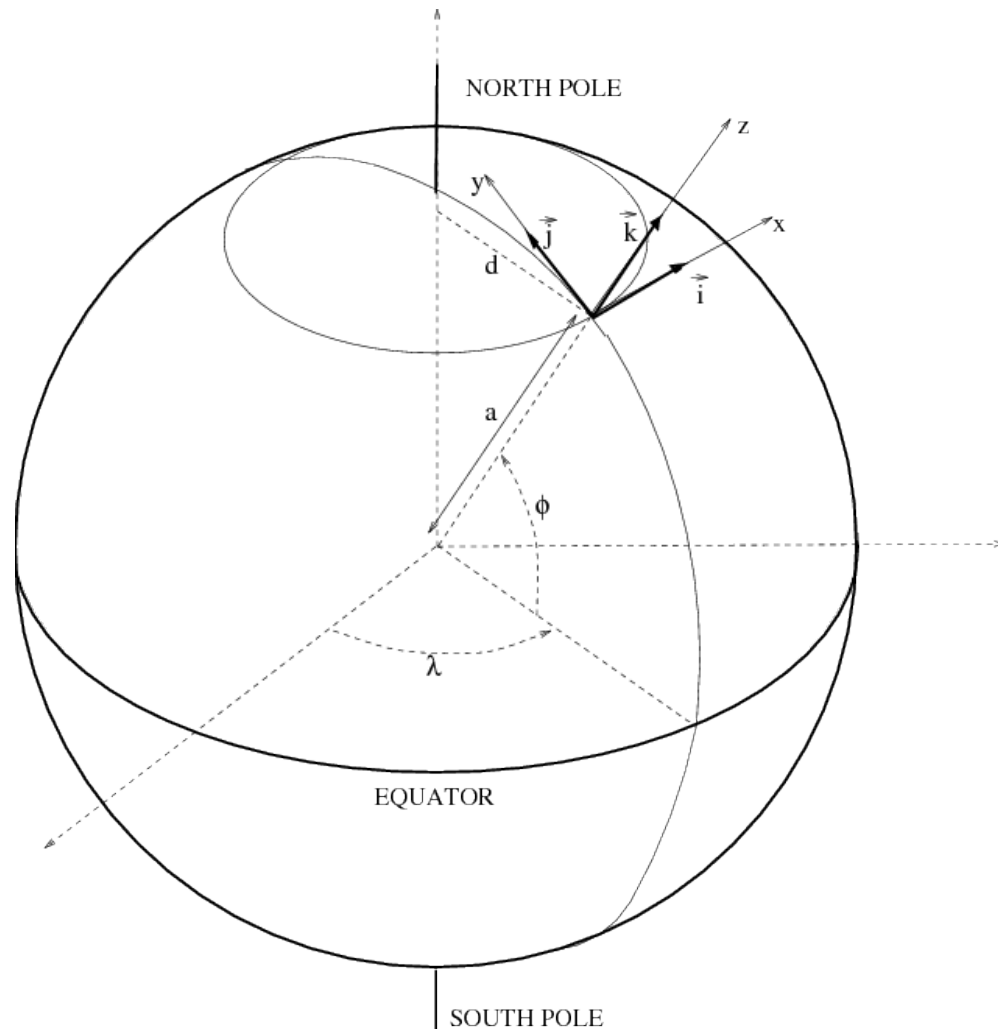
$$\theta = T \left( \frac{p_0}{p} \right)^{\kappa}$$

$\kappa = R / c_p$



## Momentum equations

Spherical coordinates, local frame



Lagrangian view  
VS  
Eulerian view

$$\frac{d}{dt} = \frac{\partial}{\partial t} + u \frac{\partial}{\partial x} + v \frac{\partial}{\partial y} + w \frac{\partial}{\partial z}$$

## Momentum equations

Momentum equations derived in the local frame

relative acceleration = pressure gradient + gravity + friction + Coriolis force + centrifugal force

$$\frac{du}{dt} - \frac{uv \tan \phi}{r} + \frac{uw}{r} = 2\Omega (v \sin \phi - w \cos \phi) - \frac{1}{\rho} \frac{\partial p}{\partial x} + Fr_x$$

$$\frac{dv}{dt} + \frac{u^2 \tan \phi}{r} + \frac{vw}{r} = -2\Omega u \sin \phi - \frac{1}{\rho} \frac{\partial p}{\partial y} + Fr_y$$

$$\frac{dw}{dt} - \frac{u^2 + v^2}{r} = 2\Omega u \cos \phi - \frac{1}{\rho} \frac{\partial p}{\partial z} - g + Fr_z$$

## Momentum equations

Approximations

thin atmosphere :  $z \ll a$

$$\frac{du}{dt} - \frac{uv \tan \phi}{a} + \frac{uw}{a} = 2\Omega \sin \phi v - 2\Omega \cos \phi w - \frac{1}{\rho} \frac{\partial p}{\partial x} + Fr_x$$

$$\frac{dv}{dt} + \frac{u^2 \tan \phi}{a} + \frac{vw}{a} = -2\Omega \sin \phi u - \frac{1}{\rho} \frac{\partial p}{\partial y} + Fr_y$$

$$\frac{dw}{dt} - \frac{u^2 + v^2}{a} = 2\Omega \cos \phi u - \frac{1}{\rho} \frac{\partial p}{\partial z} - g + Fr_z$$

- |     |                       |         |                             |
|-----|-----------------------|---------|-----------------------------|
| ●   | Hydrostatic balance   | ● ● ●   | Gradient wind balance       |
| ● ● | Geostrophic balance   | ● ● ●   | Quasi-geostrophic equations |
| ● ● | Cyclostrophic balance | ● ● ● ● | Primitive equations         |

## Momentum equations

### Thermal wind equation

Cyclostrophic balance

$$-\frac{1}{\rho} \frac{\partial p}{\partial y} = \frac{u^2 \tan \varphi}{a}$$

$\varphi$  : latitude  
 $a$  : planet radius

yields

$$2 u^2 \frac{\partial u}{\partial \zeta} = - \frac{R}{\tan \varphi} \left( \frac{\partial T}{\partial \varphi} \right)_{\zeta \text{ cst}}$$

with  $\zeta = - \ln \frac{p}{p_{\text{ref}}}$

vertical zonal wind  
 gradient

temperature  
 meridional gradient

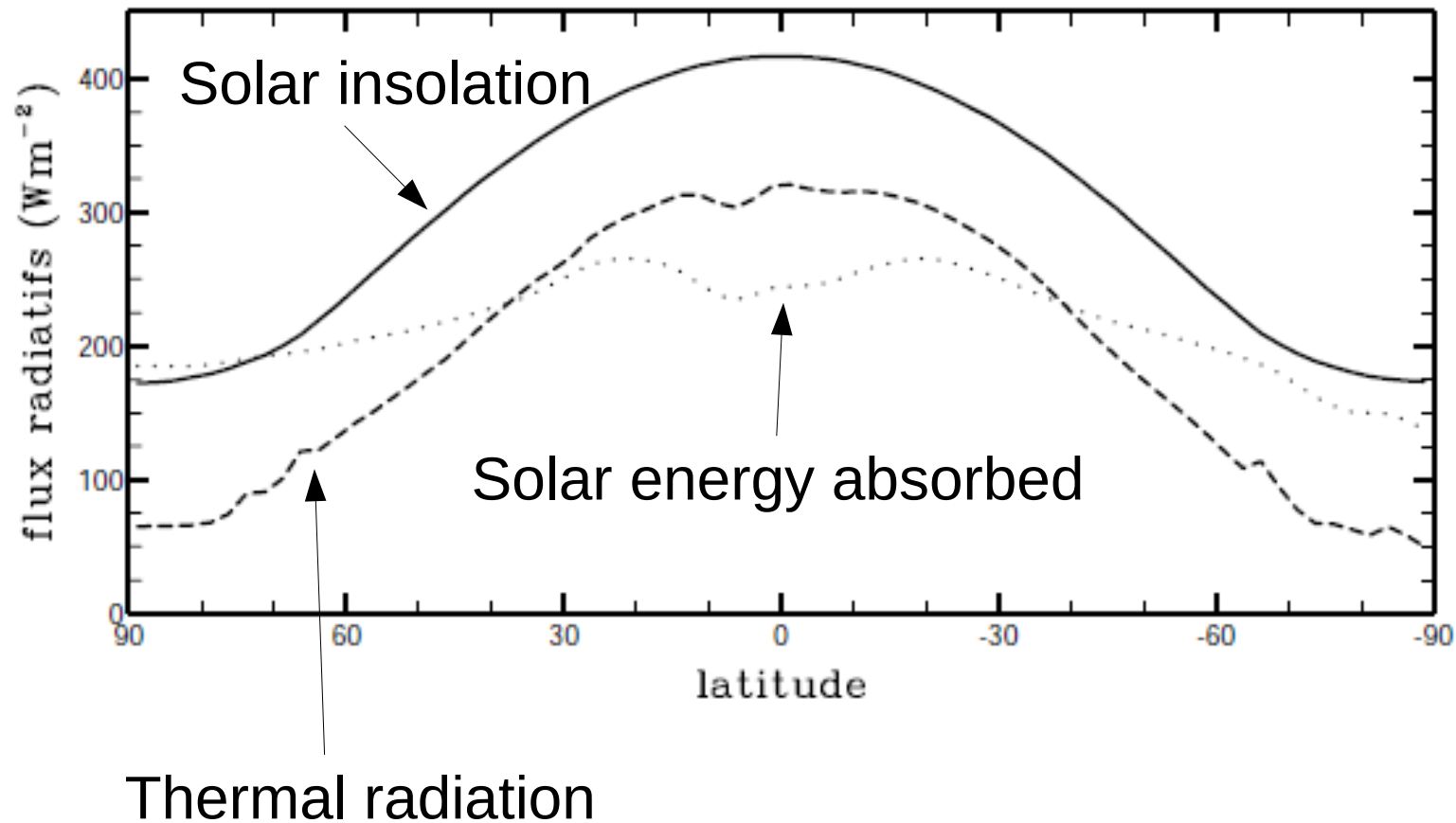
## PLANETARY ATMOSPHERES

### Atmospheric dynamics and circulation regimes

- Equations of the atmospheric fluid
- Circulation patterns in terrestrial atmospheres
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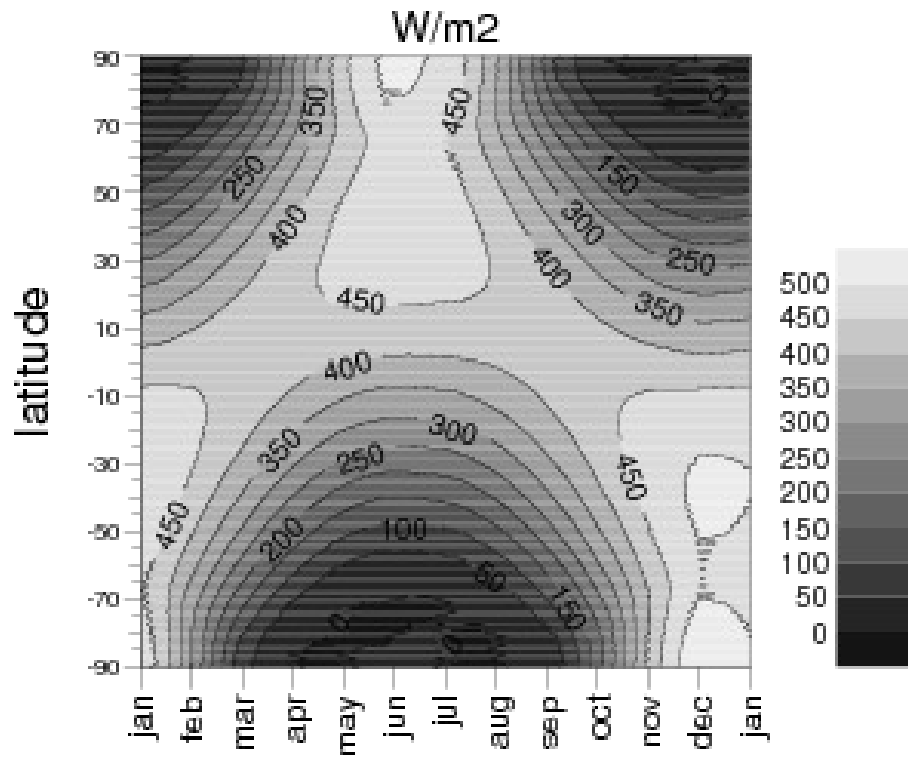
## Energy redistribution

Annual average energy balance for Earth's atmosphere

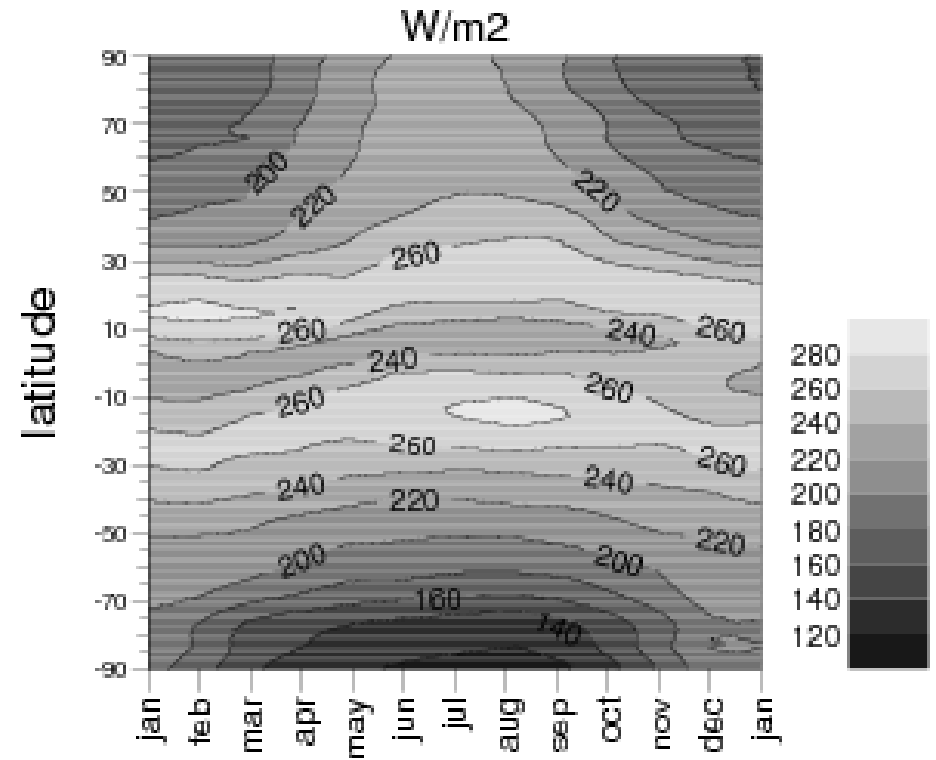


## Energy redistribution

At top of atmosphere, as a function of seasons



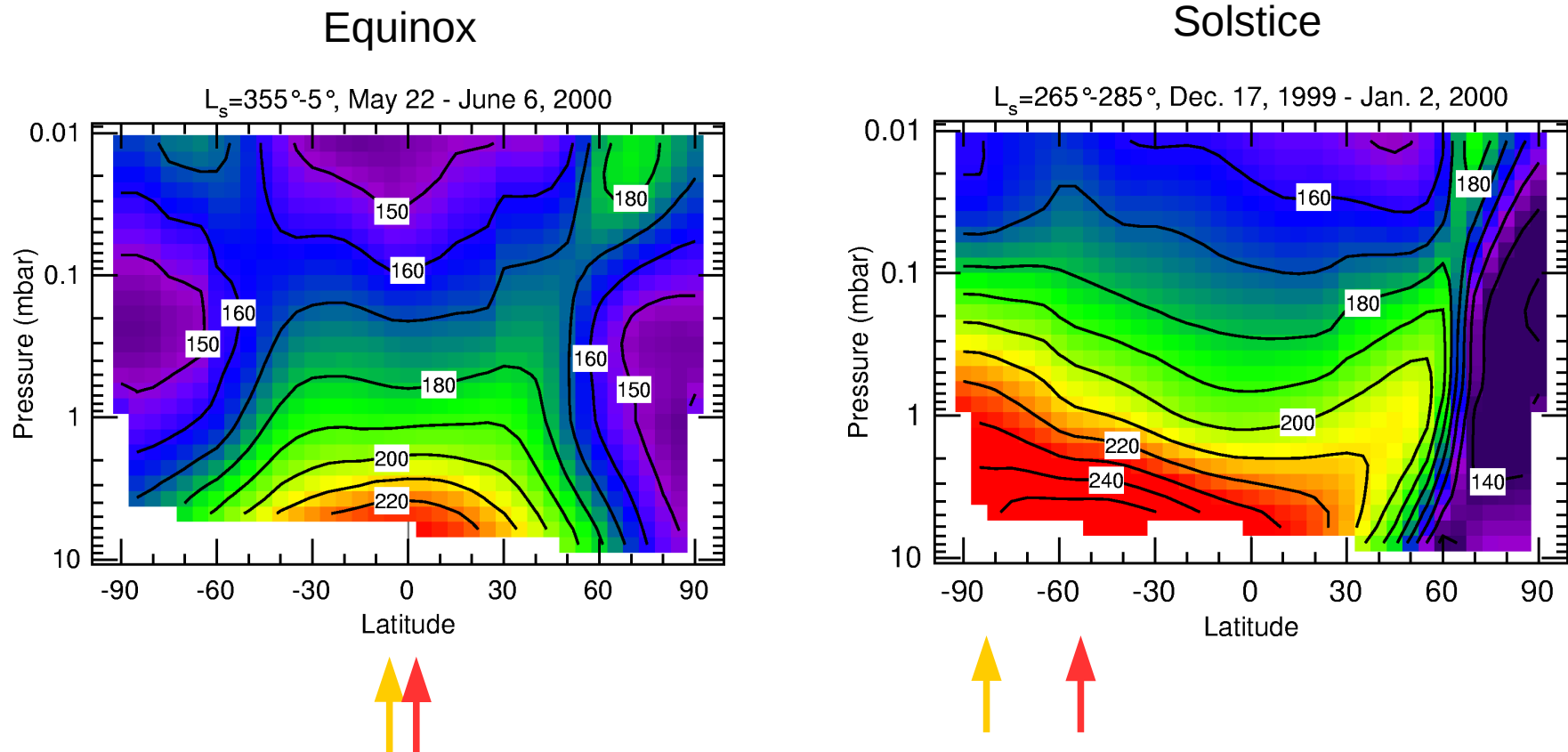
Solar insolation



Thermal radiation

## Energy redistribution

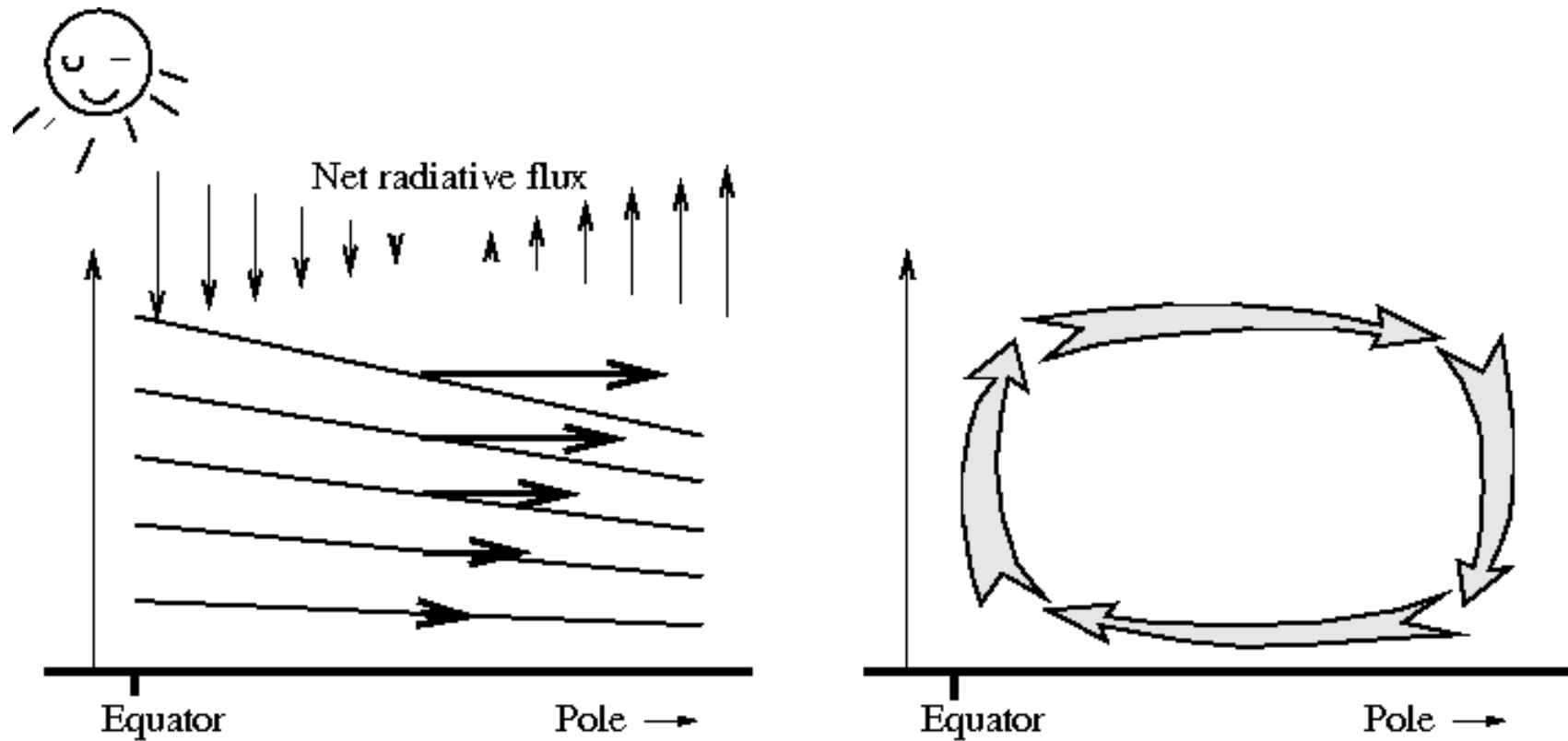
The case of Mars : role of surface thermal inertia





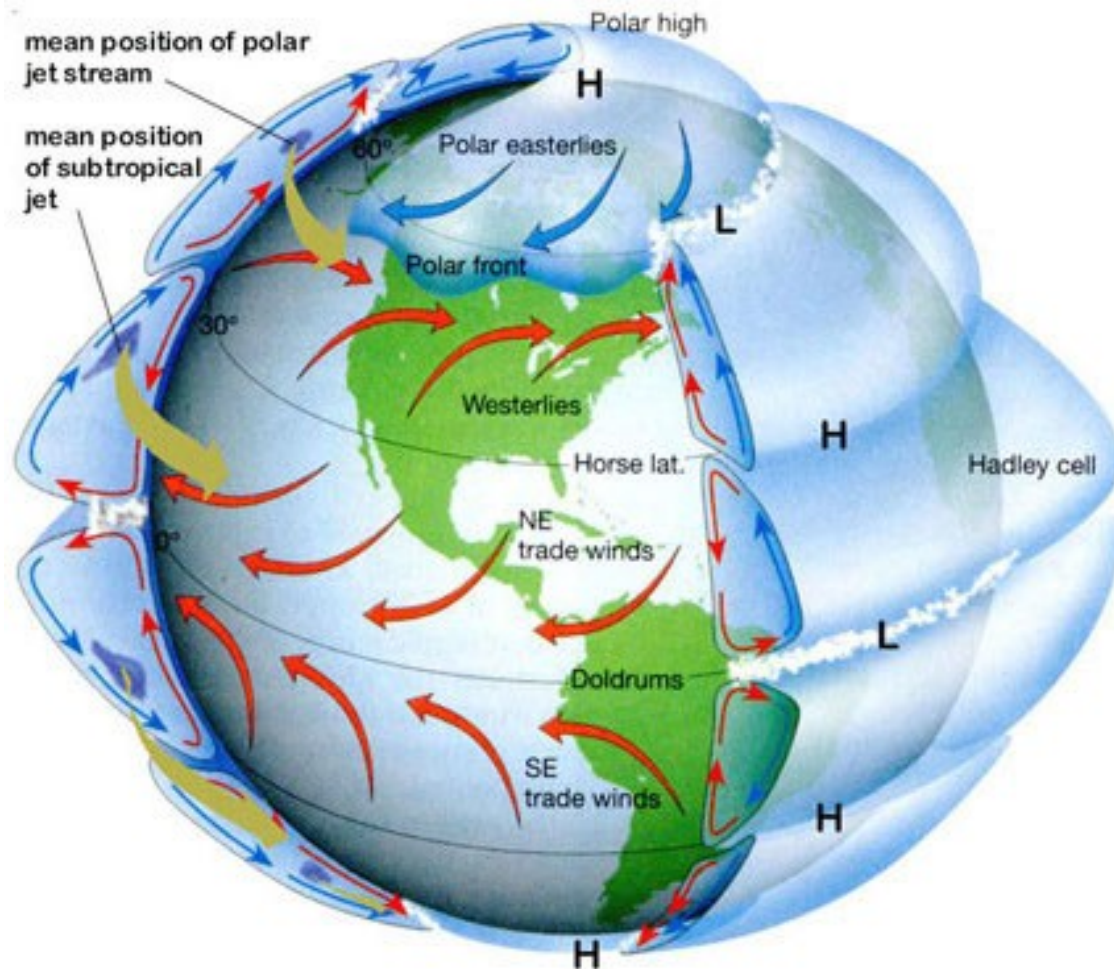
## Meridional circulation

### Driving mechanism



## Meridional circulation

Effect on the zonal wind

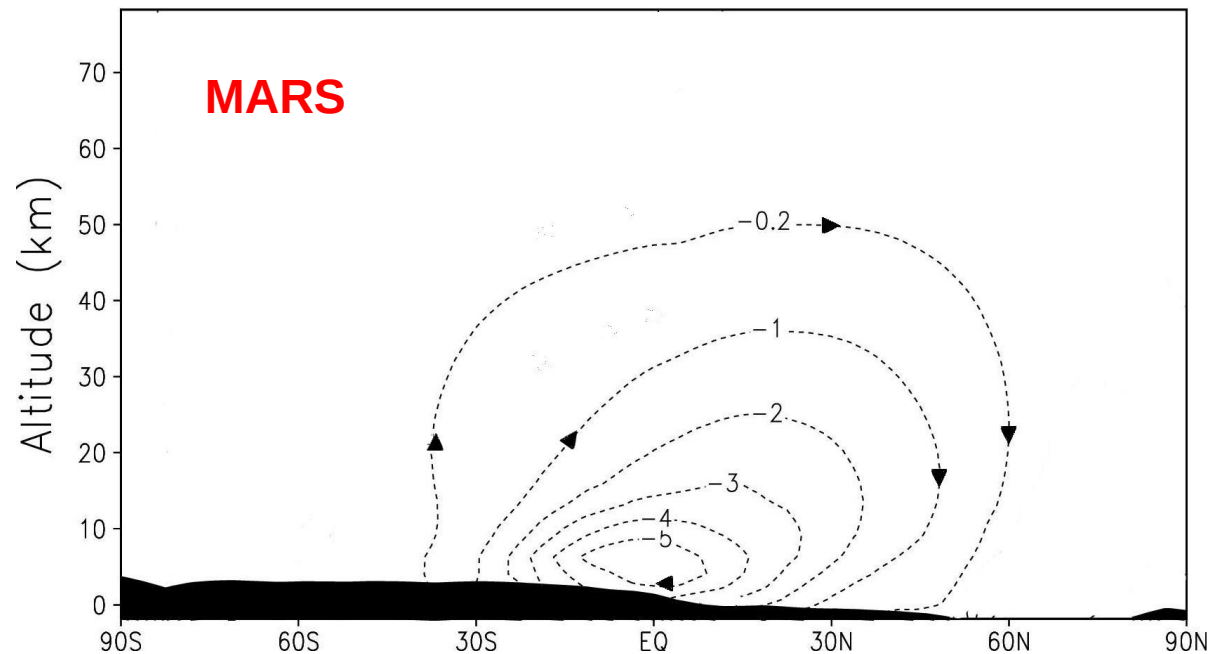
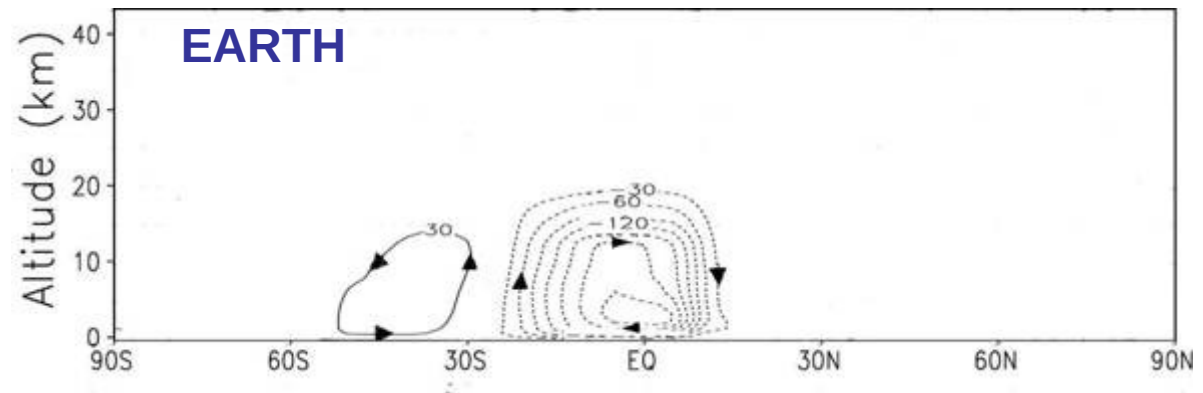


## Meridional circulation

### Mars Hadley cells

Mean meridional circulation  
(from simulations,  
mass flux in  $10^9 \text{ kg s}^{-1}$ )

Northern winter solstice

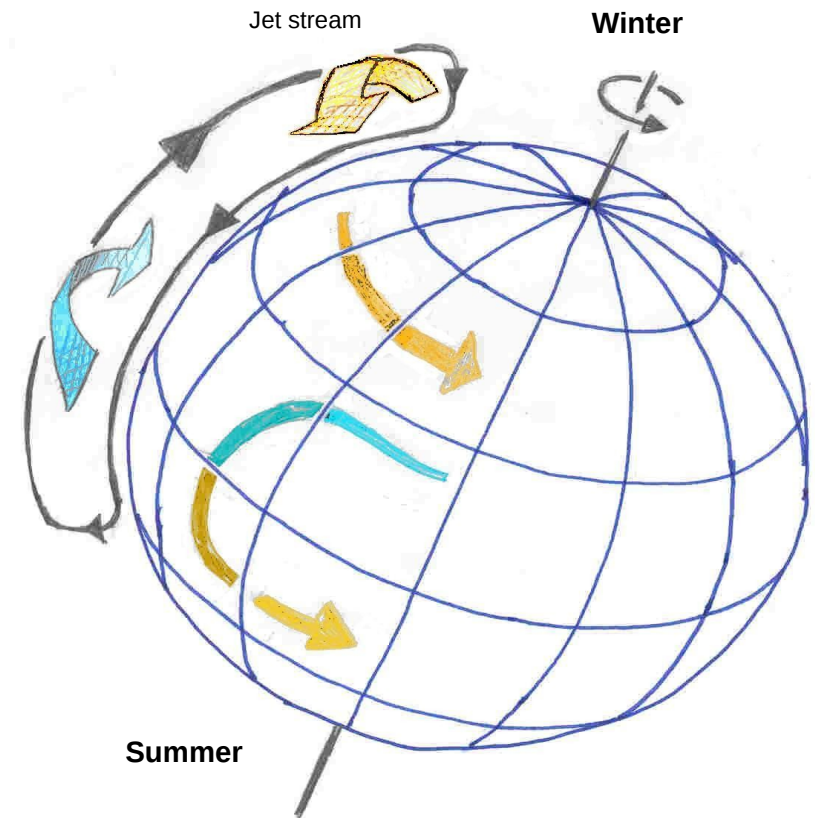
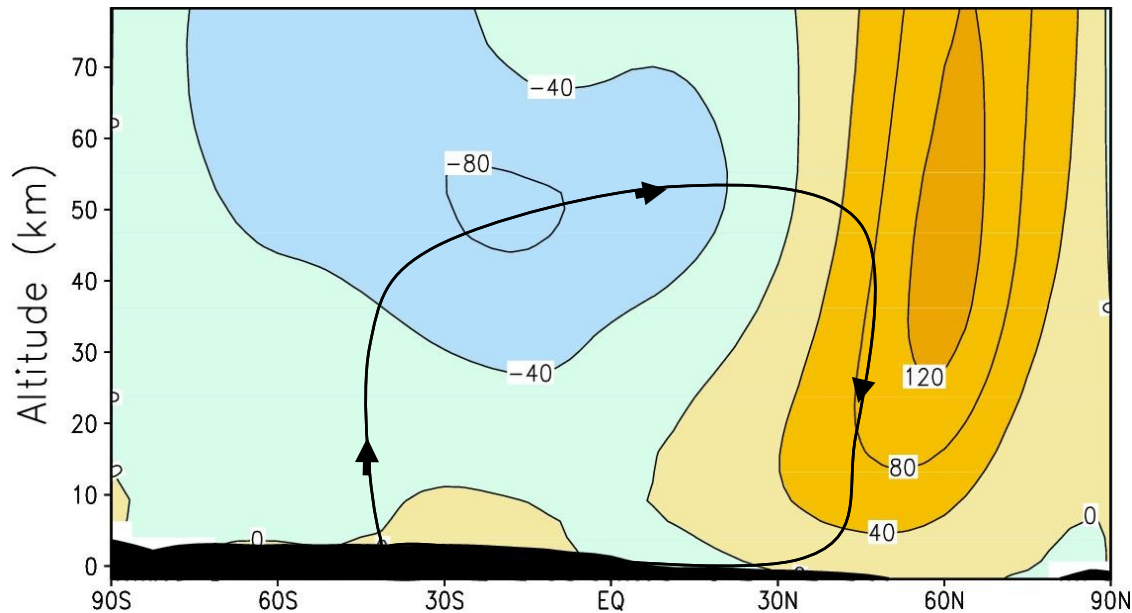


## Meridional circulation

### Mars Hadley cells

Mean zonal wind (m/s)

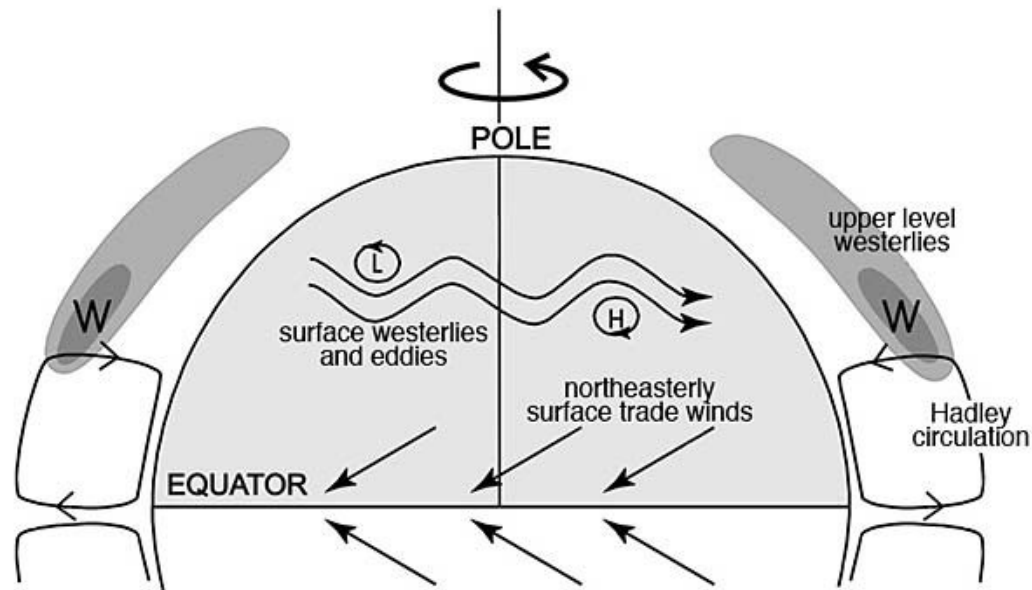
Northern winter solstice



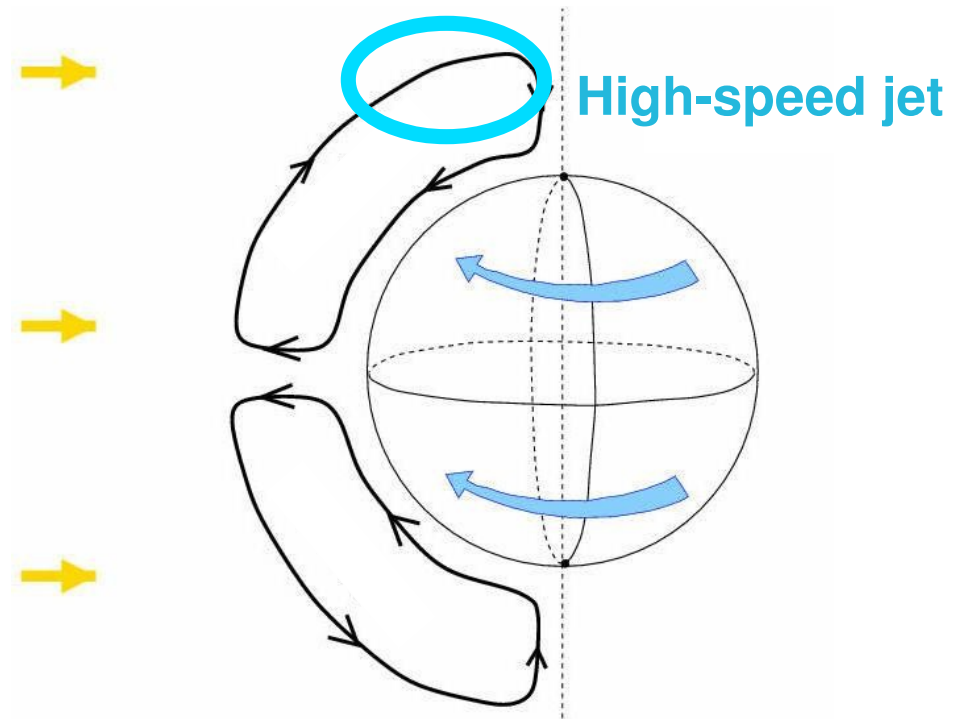
## Meridional circulation

Venus and Titan : slow rotation

Extension of Hadley cells from equator to the poles



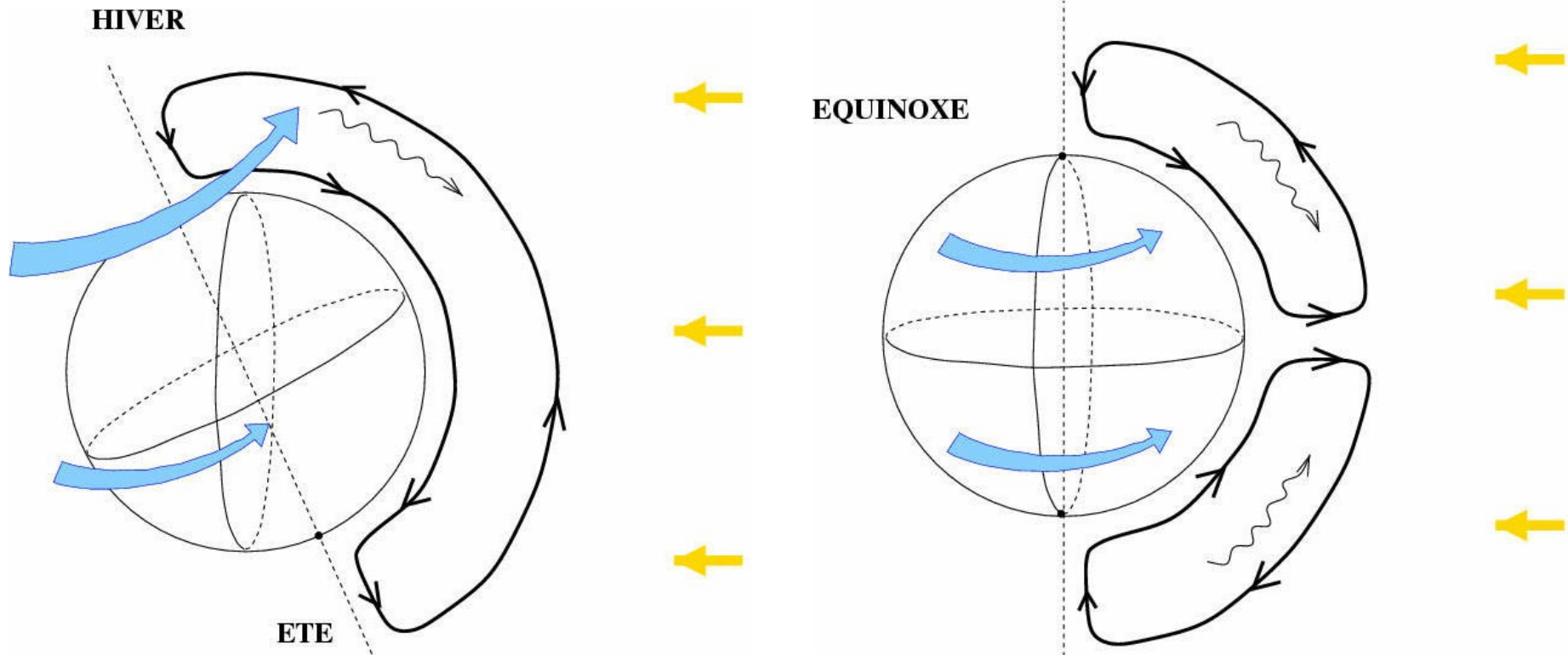
Earth



Venus

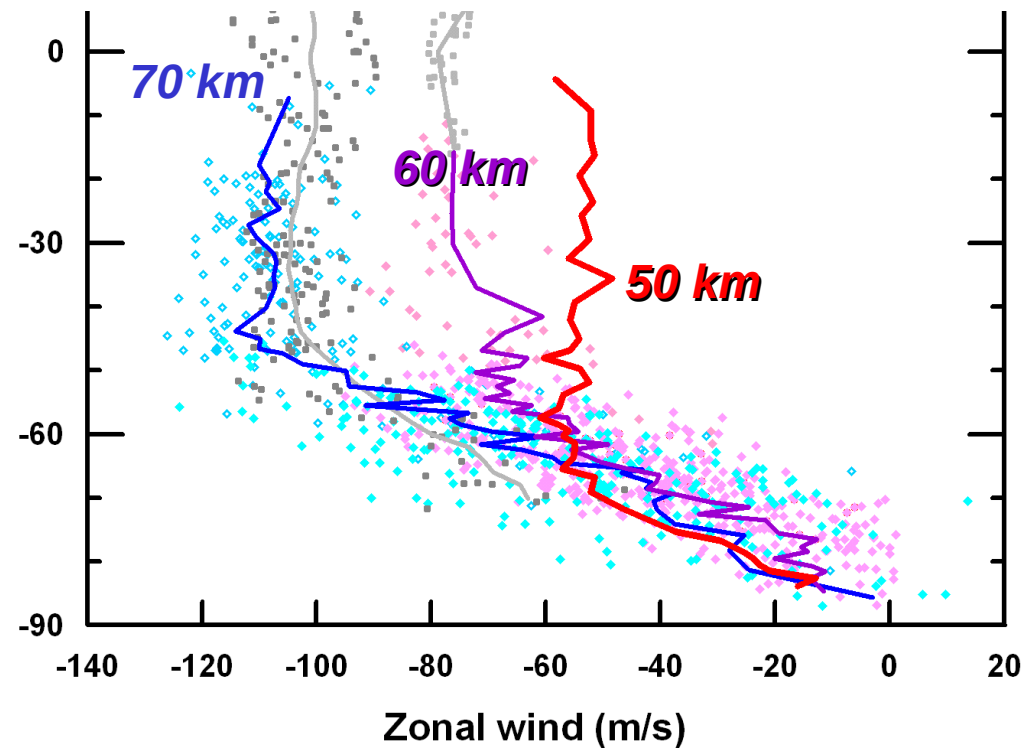
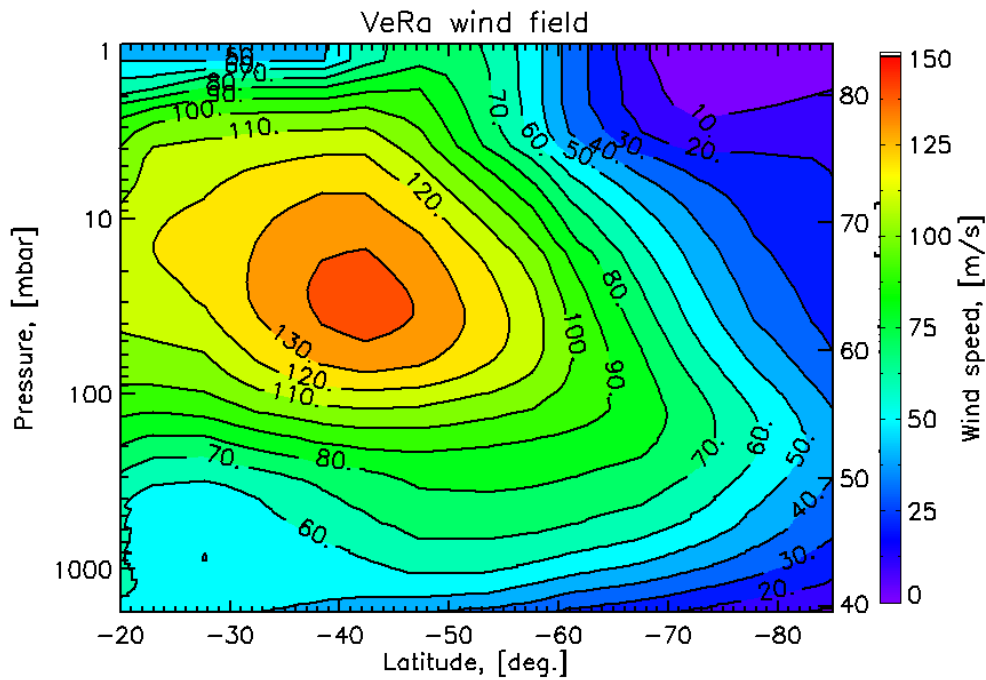
## Meridional circulation

Titan : impact of seasons



## Superrotation

### Observations : Venus

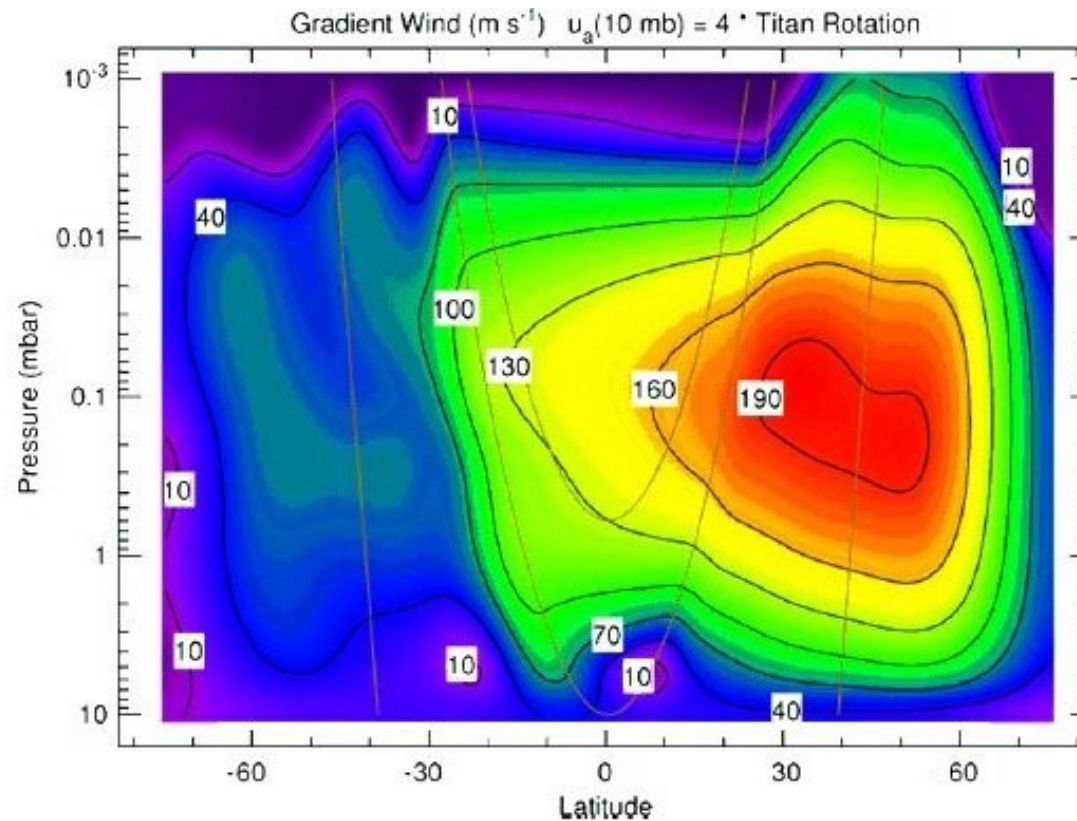


Venus Express/VeRa temperatures  
=> thermal wind equation

Venus Express/VIRTIS  
cloud tracking

## Superrotation

Observations : Titan

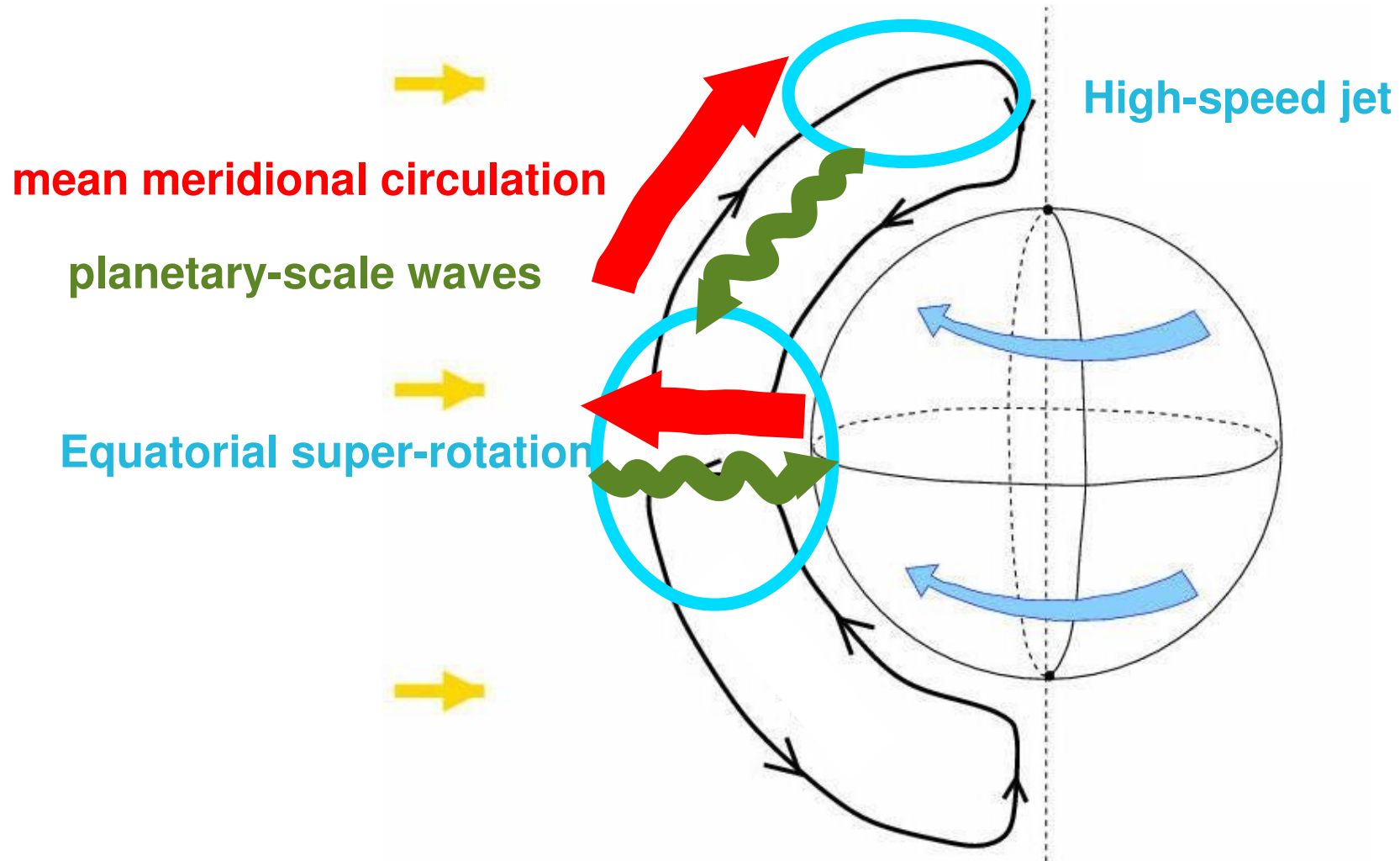


**Cassini/CIRS thermal winds retrieval  
( $L_s \sim 300^\circ$ , northern winter)**



## Superrotation

Mechanism : angular momentum transport



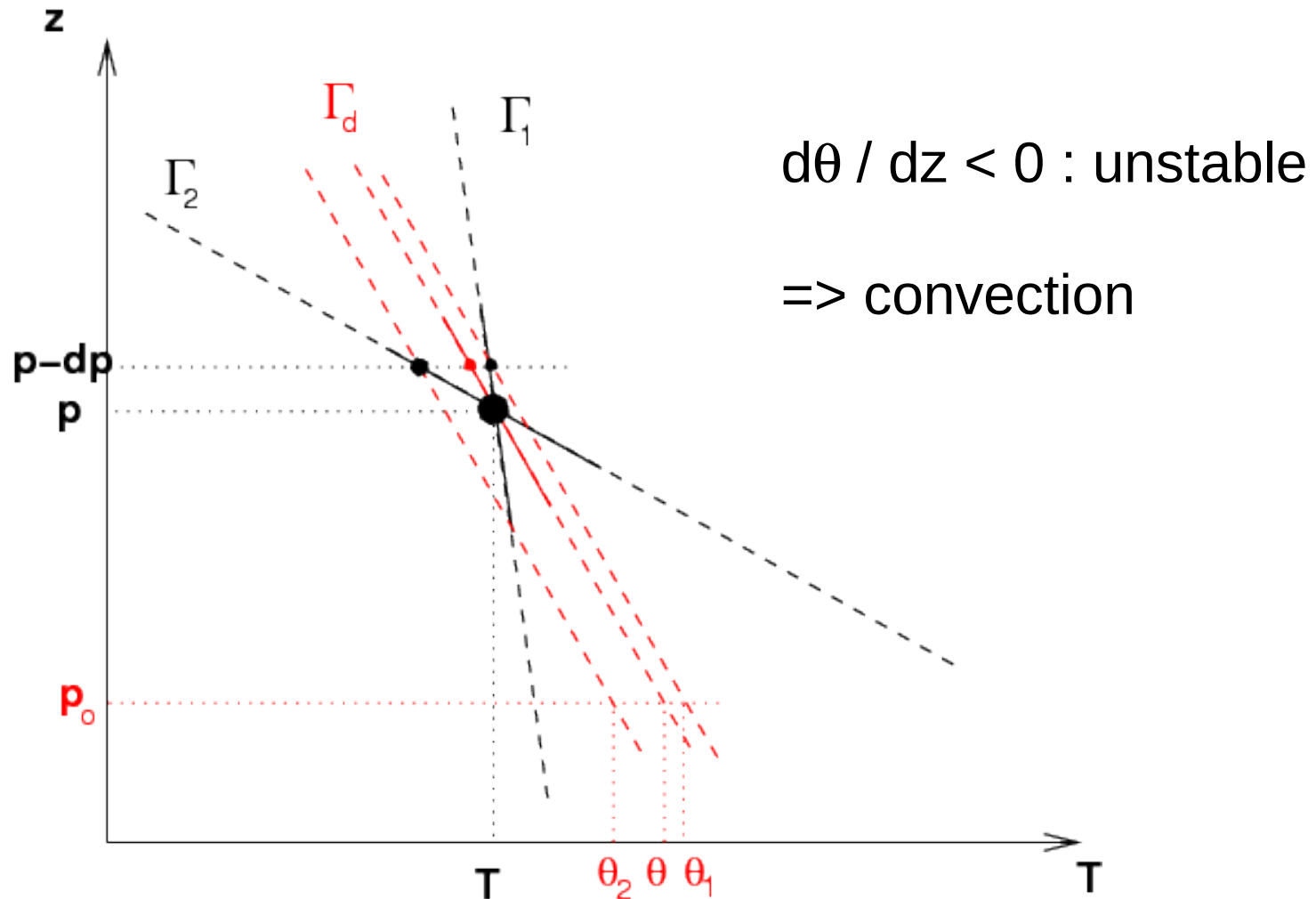
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## Convection

Buoyancy



## Convection

Static stability

$$S = \frac{dT}{dz} + \frac{g}{c_p}$$

Brunt-Väisälä

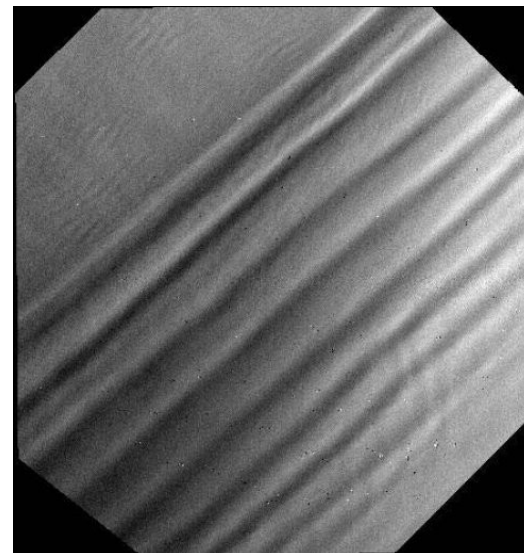
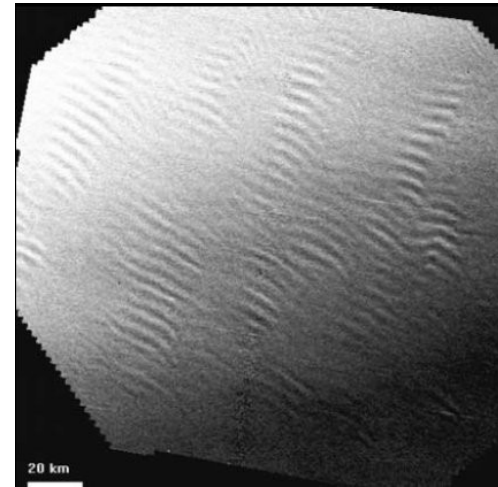
$$N_B^2 = \frac{g}{T} S = \frac{g}{\theta} \frac{d\theta}{dz}$$

Gravity waves

Mars



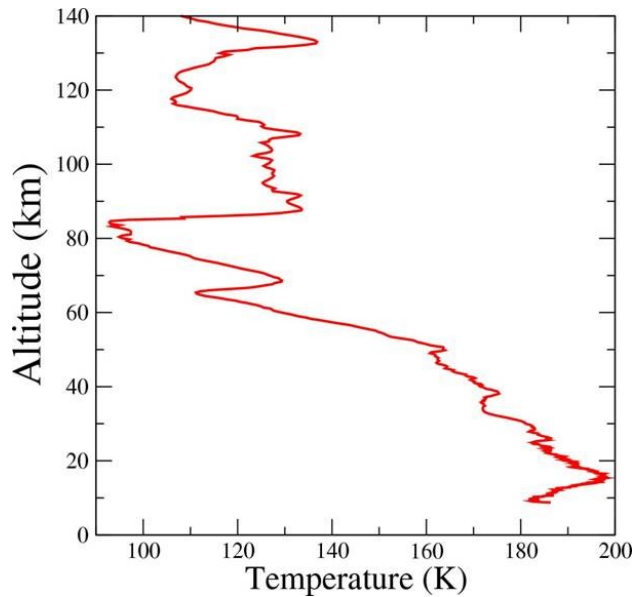
Venus



## Gravity waves

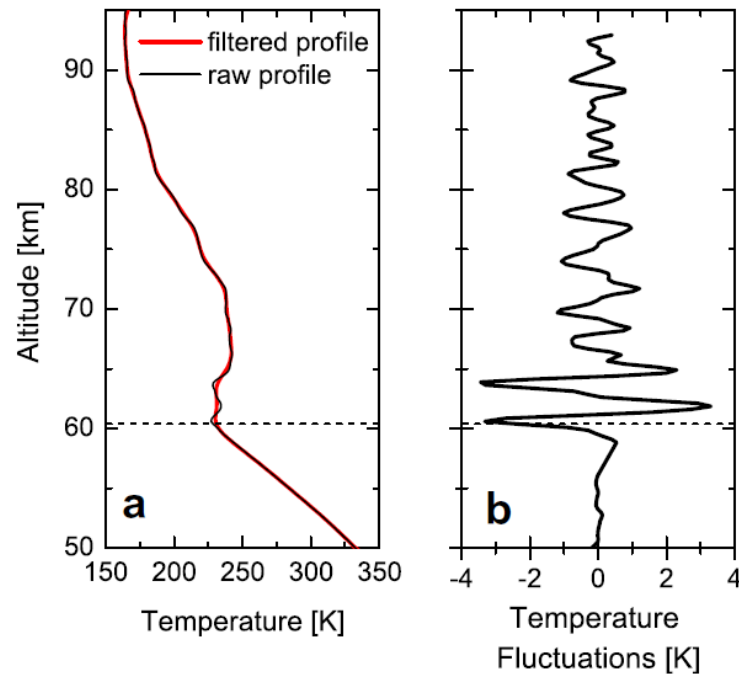
### Mars

#### Pathfinder entry profile



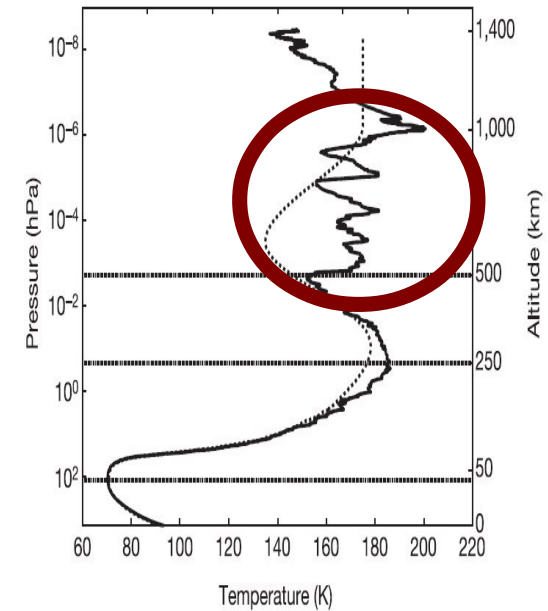
### Venus

#### VenusExpress/VeRa



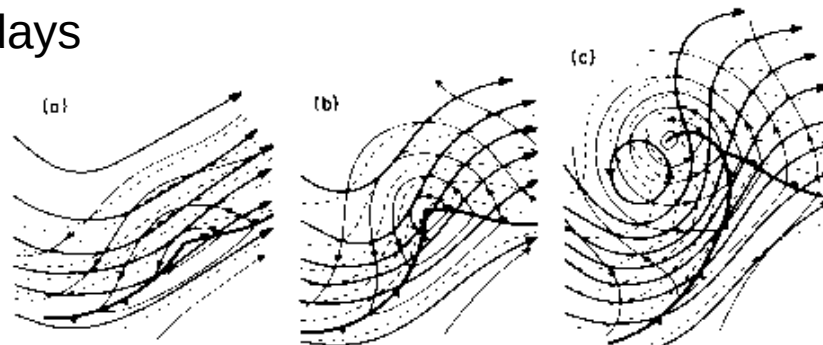
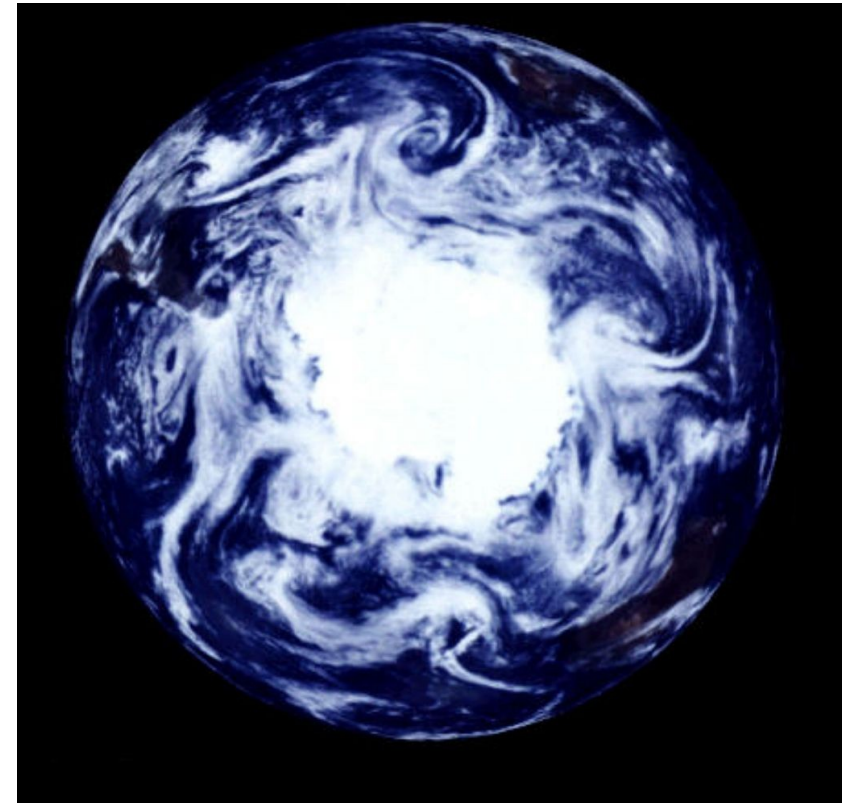
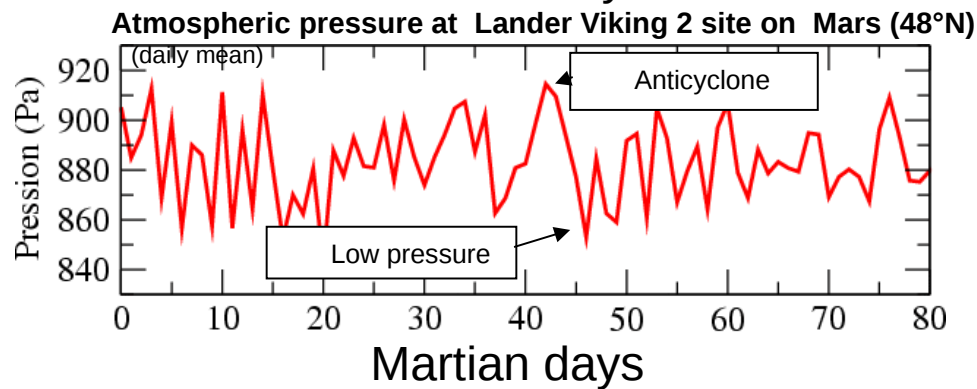
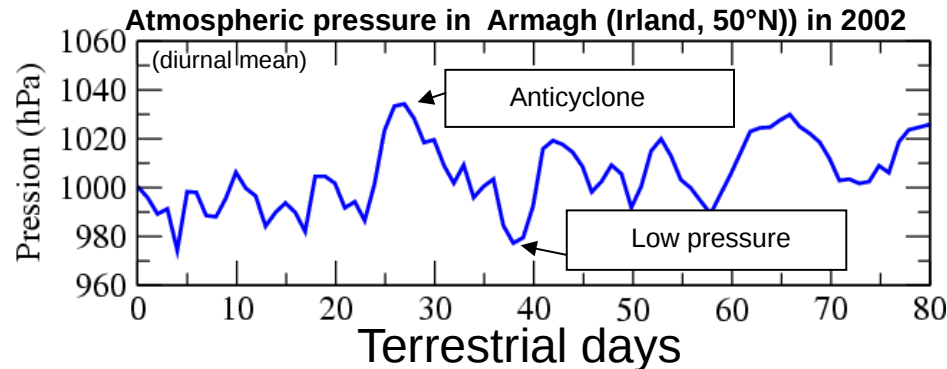
### Titan

#### Huygens/HASI



## Planetary-scale waves

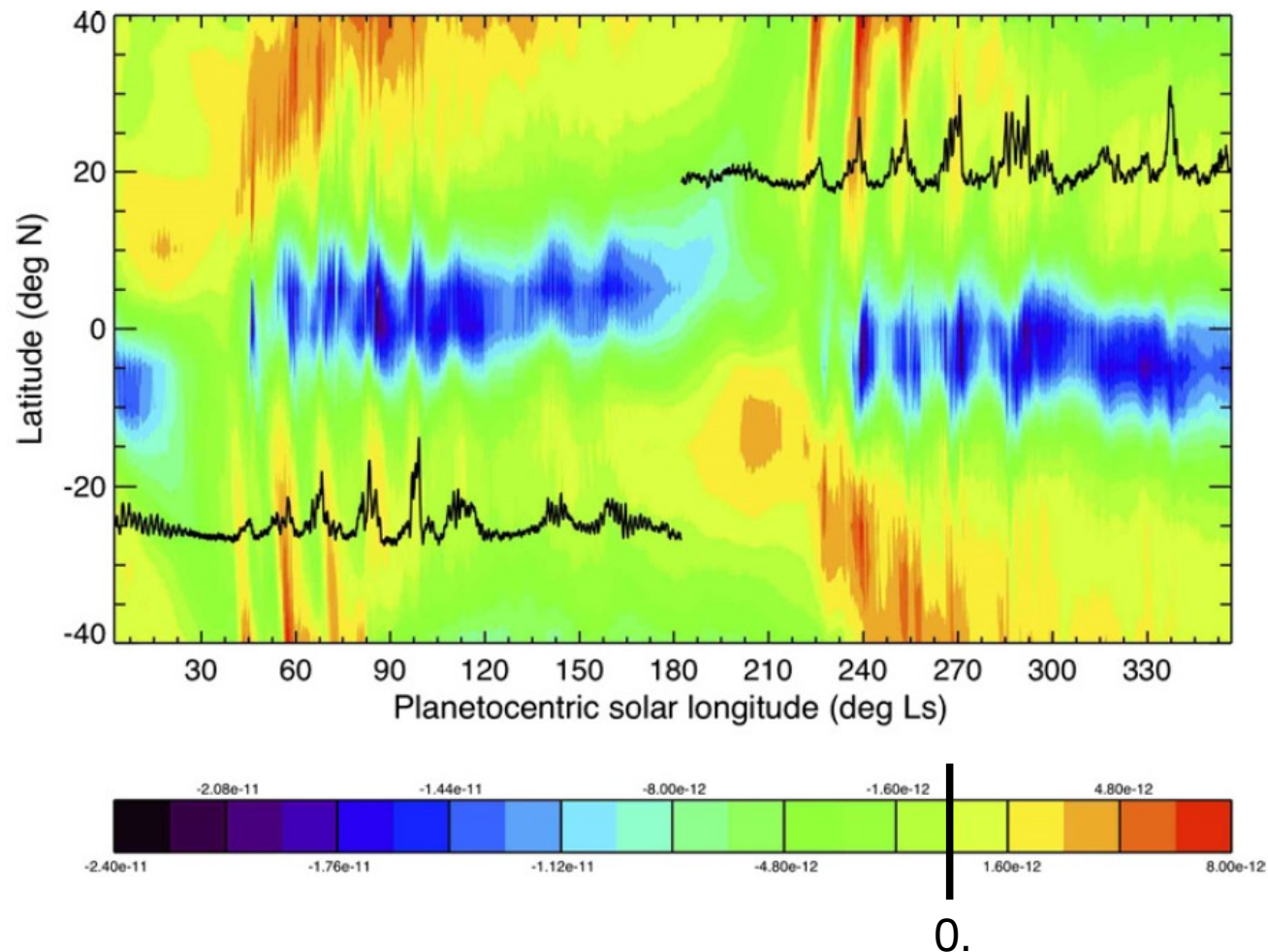
### Baroclinic waves



## Planetary-scale waves

### Barotropic waves

Barotropic instability criterion (colors)  
and angular momentum transport by waves (black line)  
in a climate model of Titan's stratosphere





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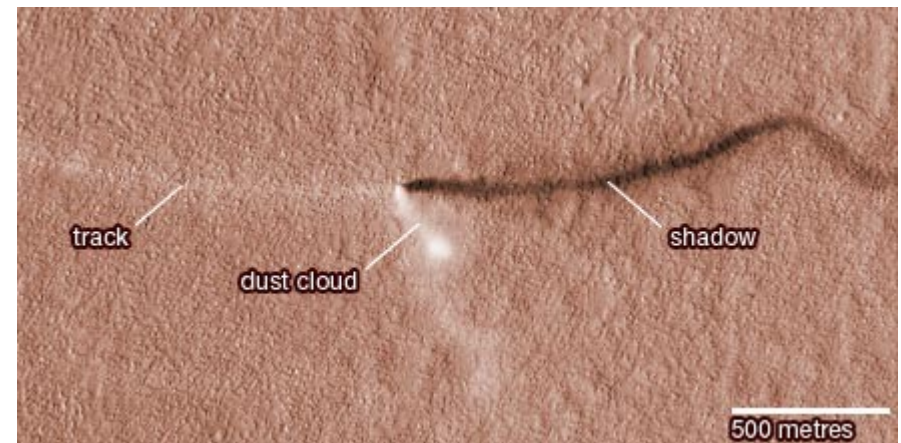
## Small-scale vortices

Dust devils

EARTH



MARS



## Small-scale vortices

### Tornadoes



## Synoptic vortices

### Earth hurricanes



### Earth extra-tropical cyclones



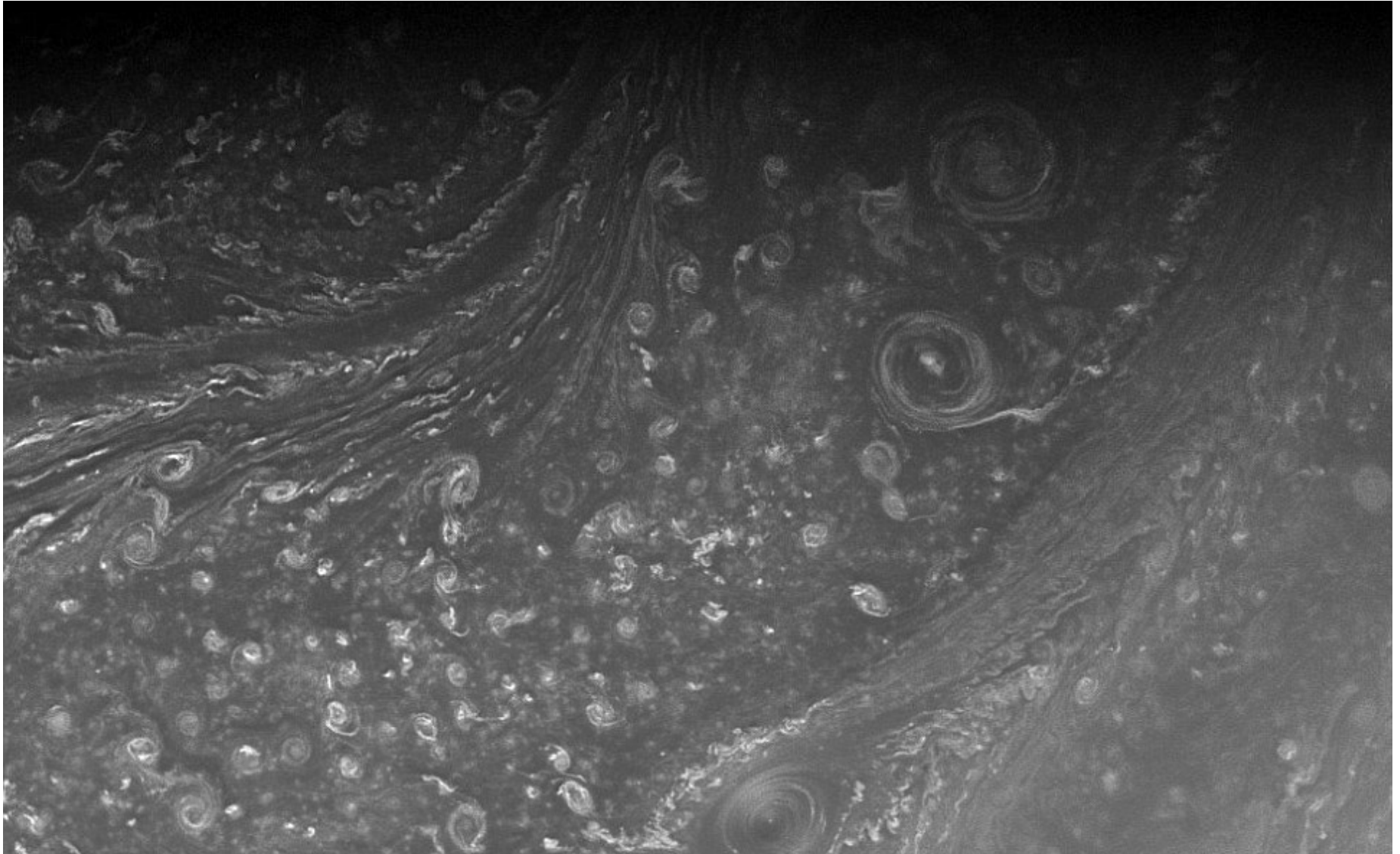
## Giant planet vortices

Jupiter (Voyager 1)

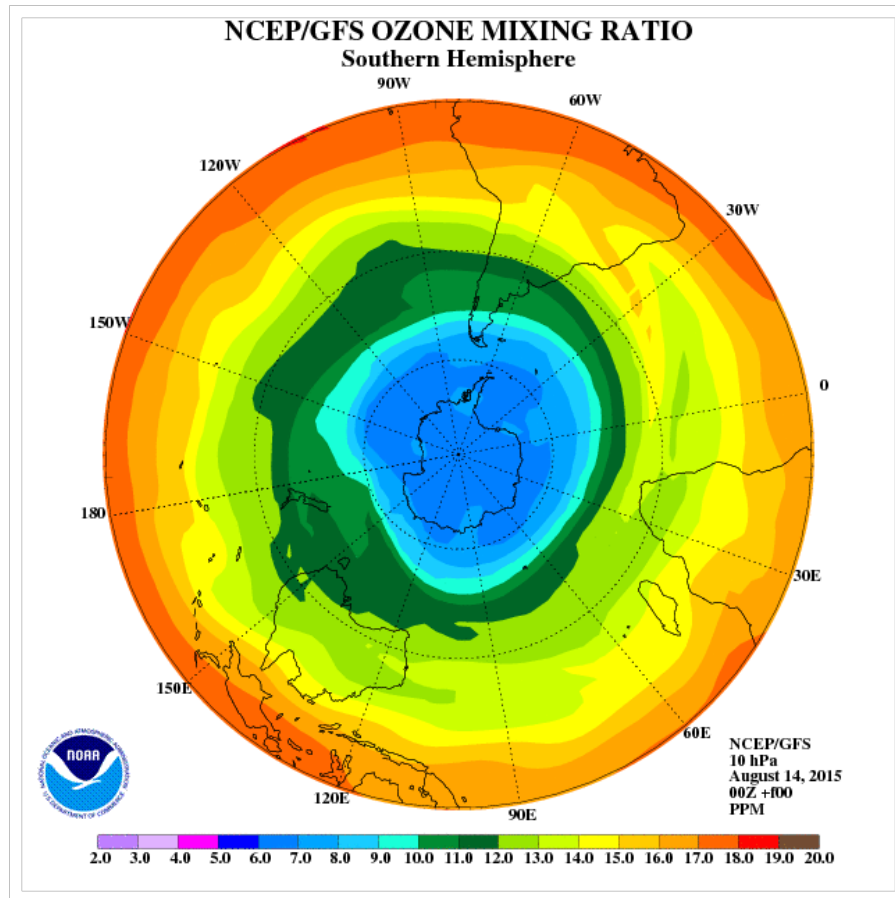


## Giant planet vortices

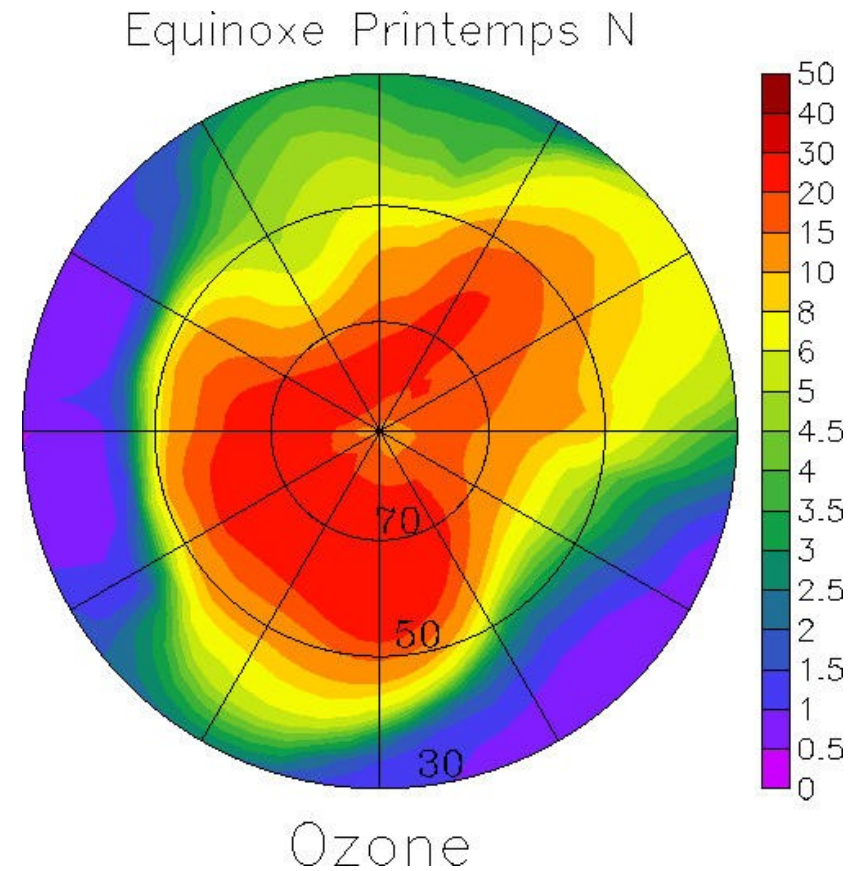
Saturn (Cassini/ISS)



## Polar vortices



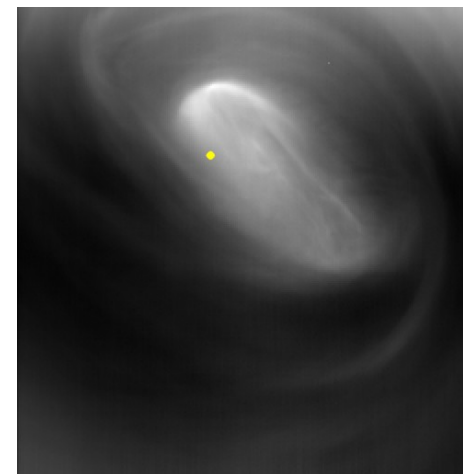
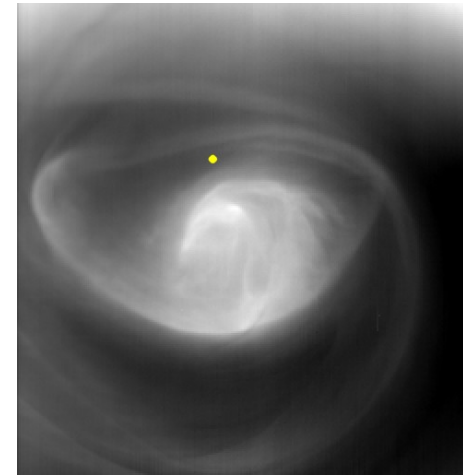
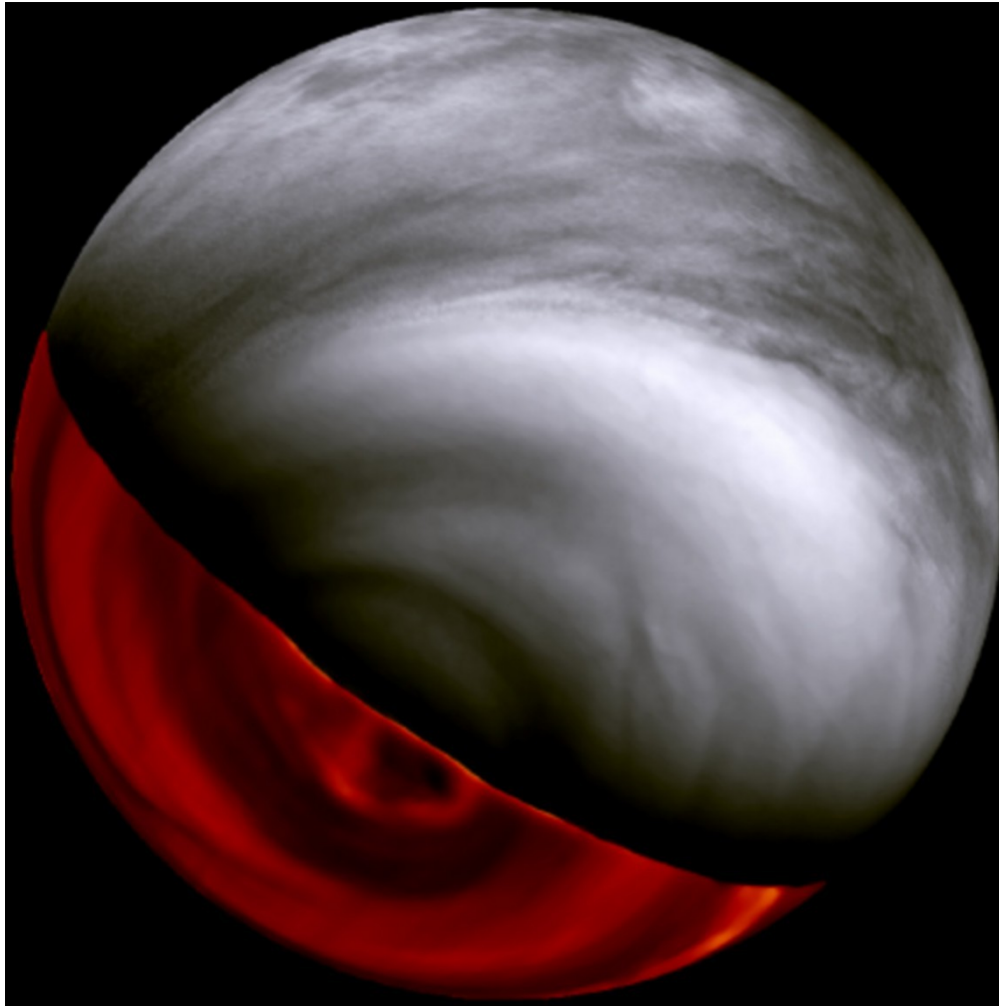
**EARTH (observation)**



**MARS (model)**

## Polar vortices

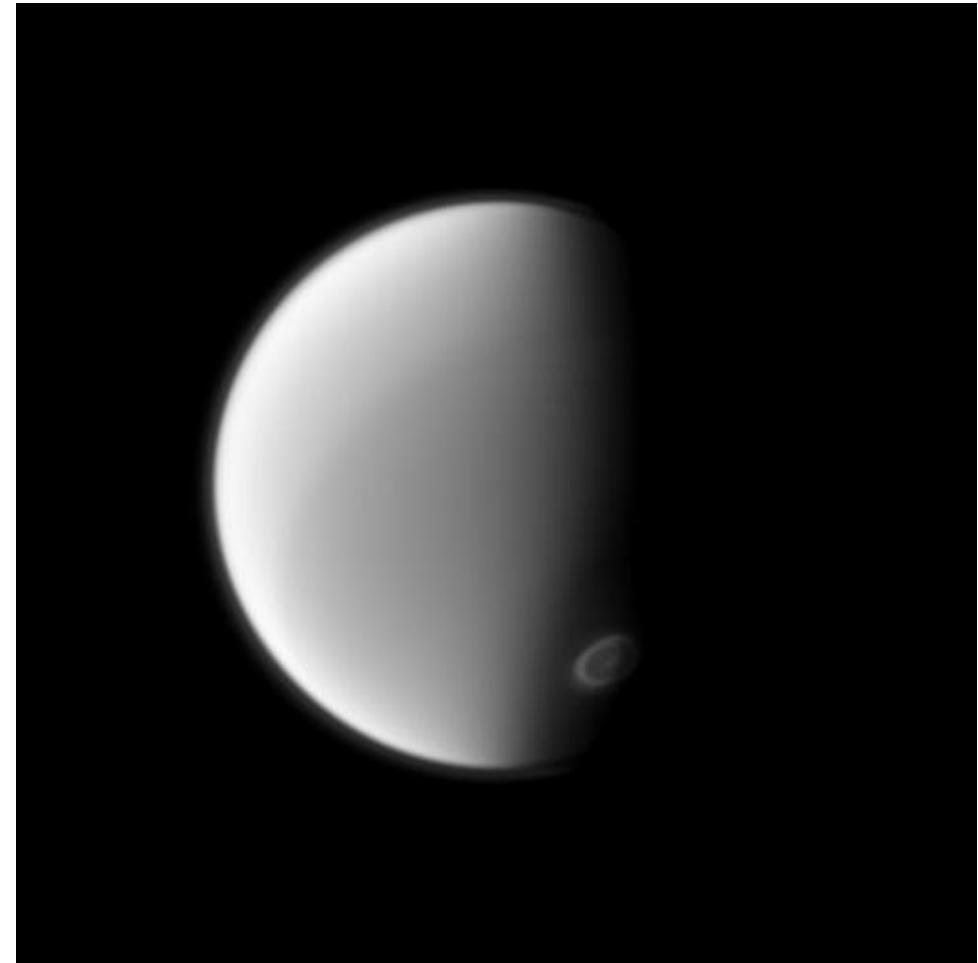
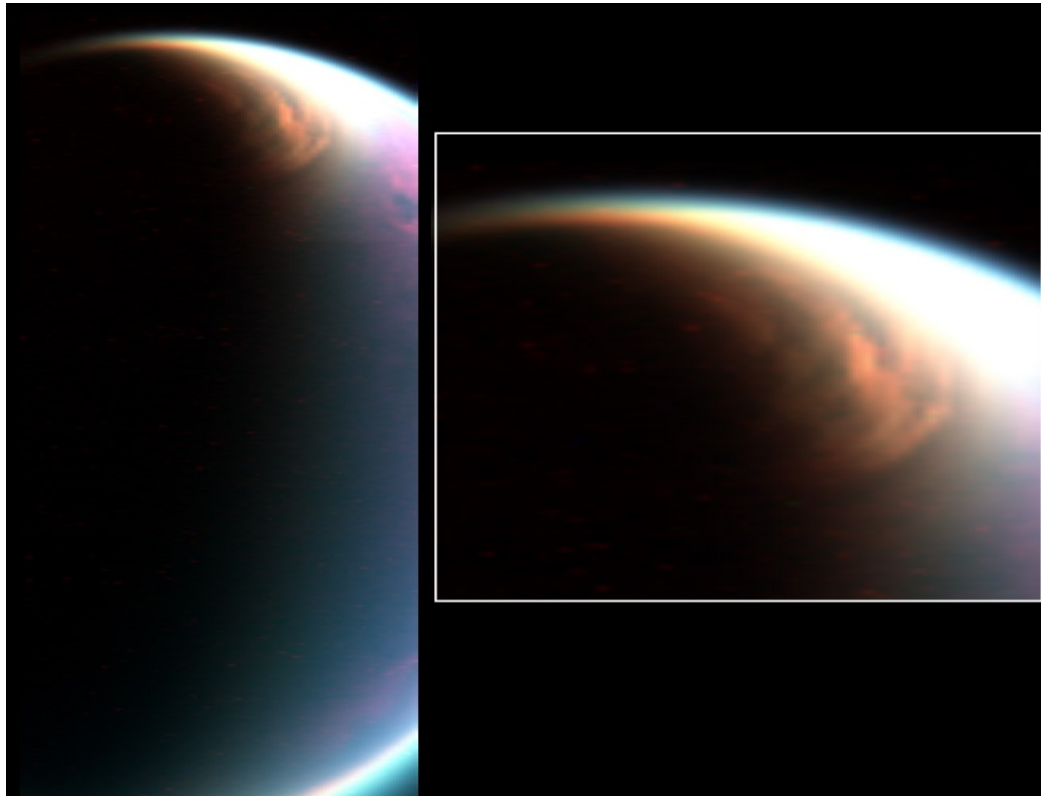
VENUS





## Polar vortices

TITAN



**Polar vortices**

Saturn (Cassini/ISS)

